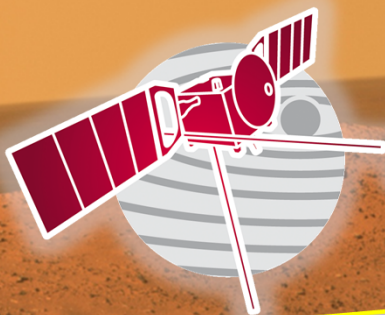


Atmospheric aerosols properties via solar infrared occultation observations by SPICAM IR



mars express



**Scientific Workshop:
“From Mars
Express to
ExoMars”**

**27–28 February 2018,
ESAC Madrid, Spain**



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SPICAM IR

Infrared spectrometer

- Spectral range : 1–1.7 μm
- Spectral resolution : $\approx 3.5\text{--}4\text{ cm}^{-1}$
- Employing an Acousto-Optic Tunable Filter (AOTF)
- Resolving power :
from 1800 at 1.6 μm to 2400 at 1.1 μm
- Objects of interest :
gaseous CO_2 (absorption bands in the range of 1.4–1.65 μm), O_2
gaseous H_2O (in the 1.37 μm band),
aerosols (mineral dust and H_2O ice).
- Launch - July 2003,
- Start of scientific operations – 2004



Spacecraft to limb distance

from 1000 to 13000 km

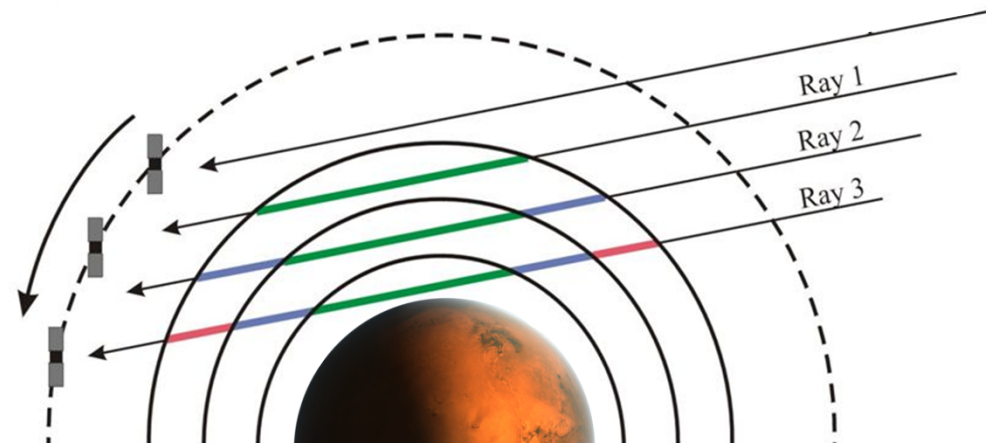
Vertical resolution

4–5 km (varies from 1 to 12 km)

Field of view

4.2 arc min (1.2 mrad)

Solar Occultation



Data retrieval

Beer–Lambert–Bouguer law :

$$I(l) = I_0 e^{-k_\lambda l}$$

$$\tau_\lambda(L) = 2 \int_{h_0}^{\infty} k_\lambda(l) dl$$

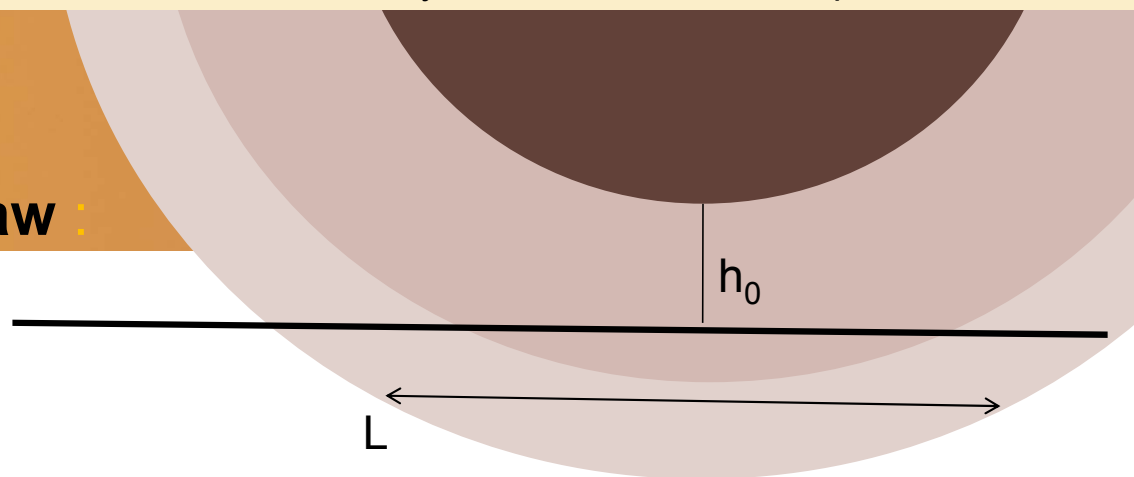
$$\tau_\lambda(L) = -\ln(I_\lambda(L)/I_0)$$

At each altitude layer:

$$k_{\text{ext}}(\lambda, z) = \int_0^{\infty} \sigma_{\text{ext}}(r, z) n(r, z) dr$$

Size distribution:

$$n(r) = \text{const} \times r^{-1} \exp\left(-\frac{(\ln r - \ln r_g)^2}{2 \ln^2 \sigma_g}\right)$$

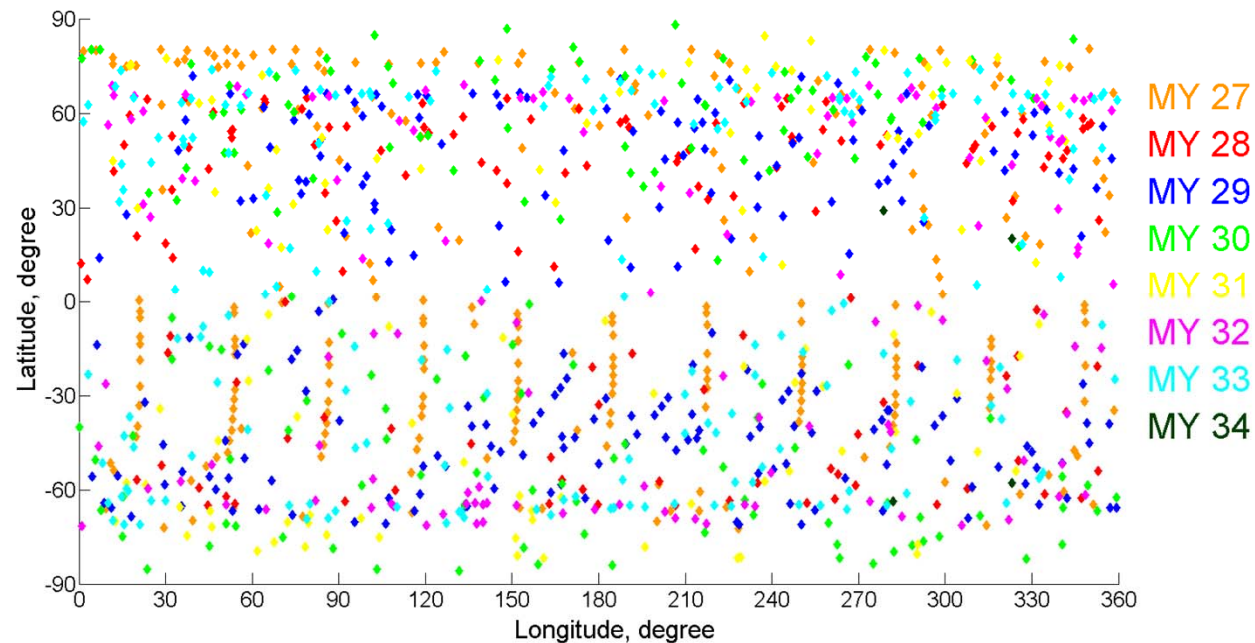
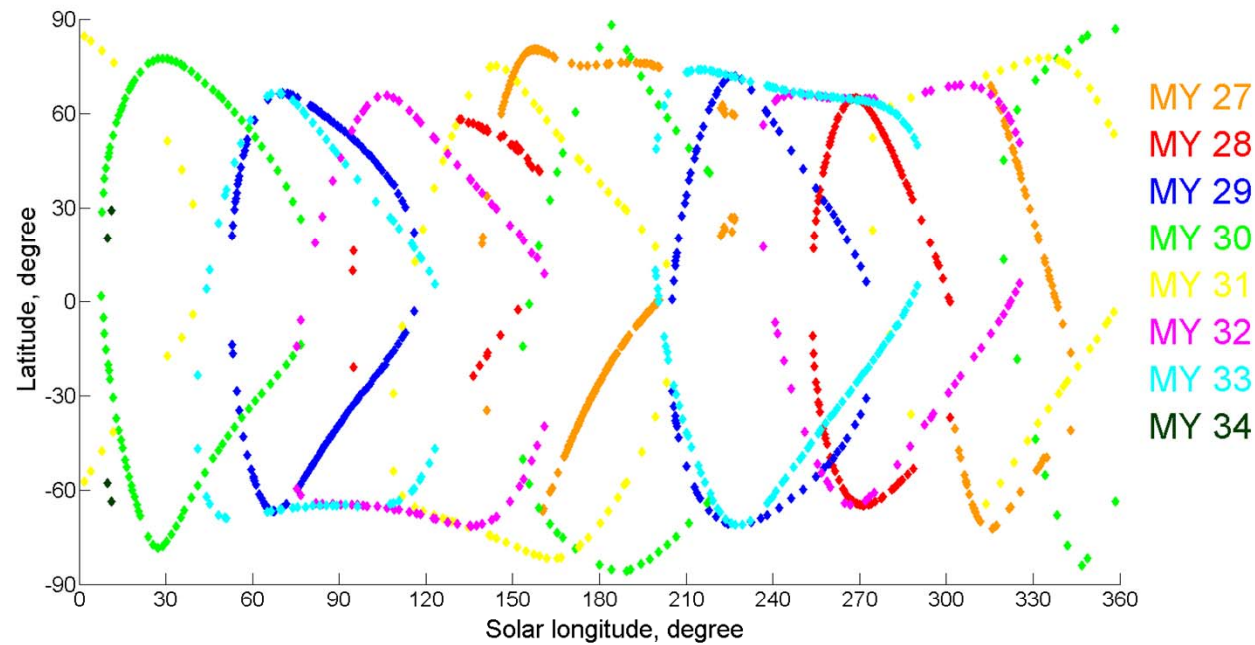


$$r_{\text{eff}} = r_g \exp\left(\frac{5}{2} \ln^2 \sigma_g\right)$$

$$v_{\text{eff}} = \exp(\ln^2 \sigma_g) - 1$$

Number density:

$$N(z) = \frac{k_{\text{ext}}(z)}{\int_0^{\infty} \pi r^2 Q_{\text{ext}}(\lambda, z) n(r, z) dr}$$



Coverage and distribution of observations

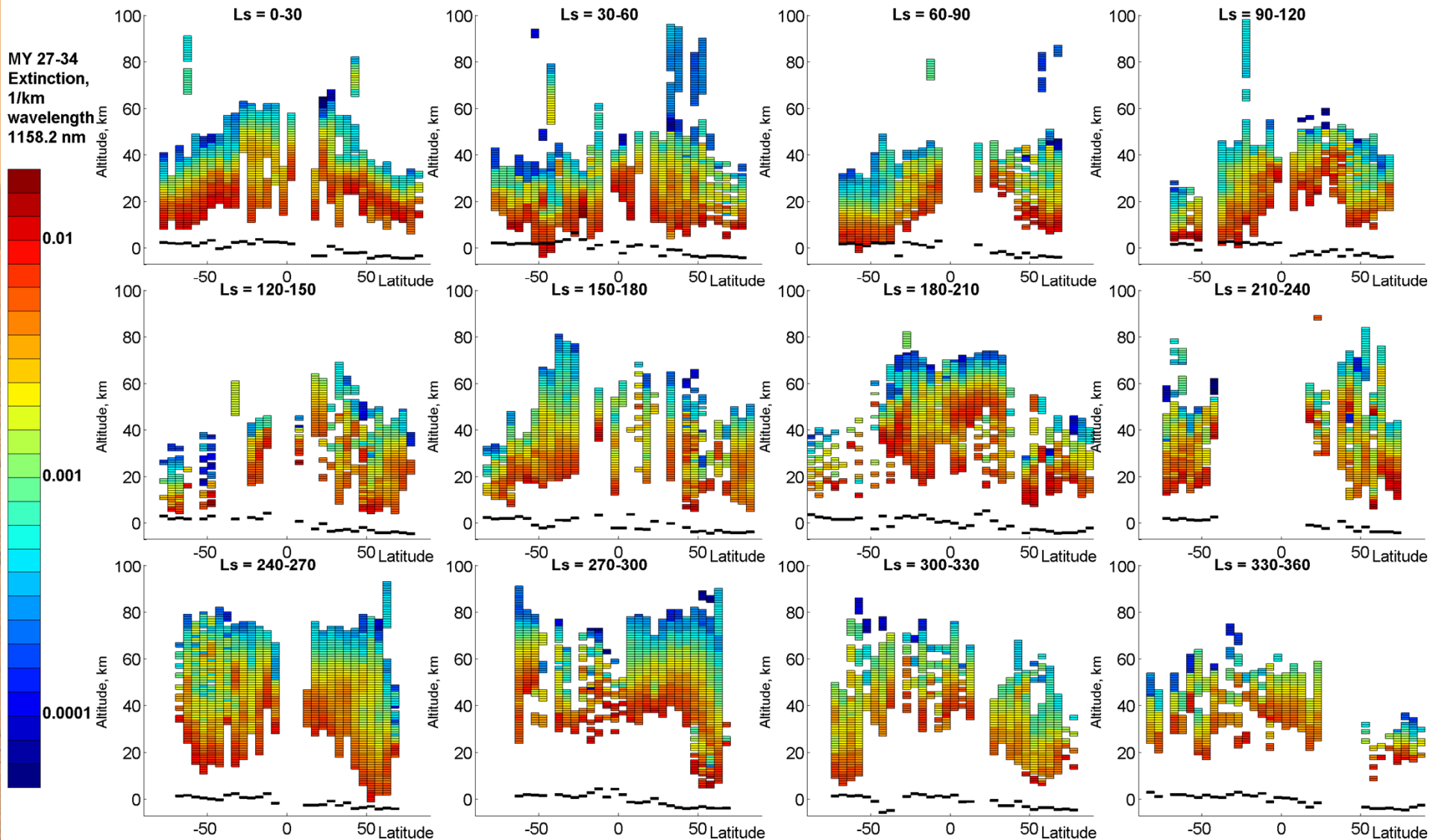
- Seasonal

- Spatial

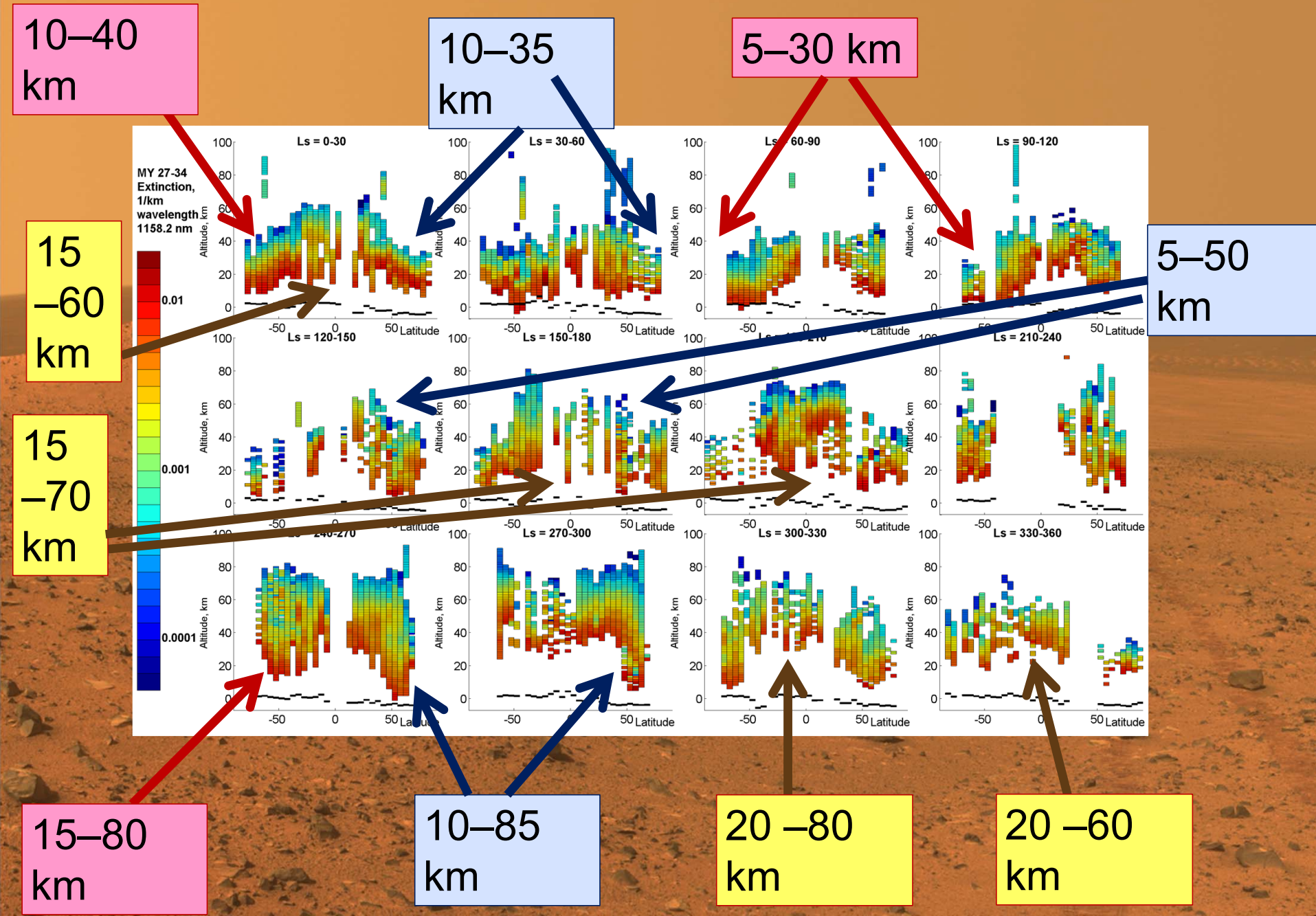
Since January 2005
– till now

Extinction coefficient

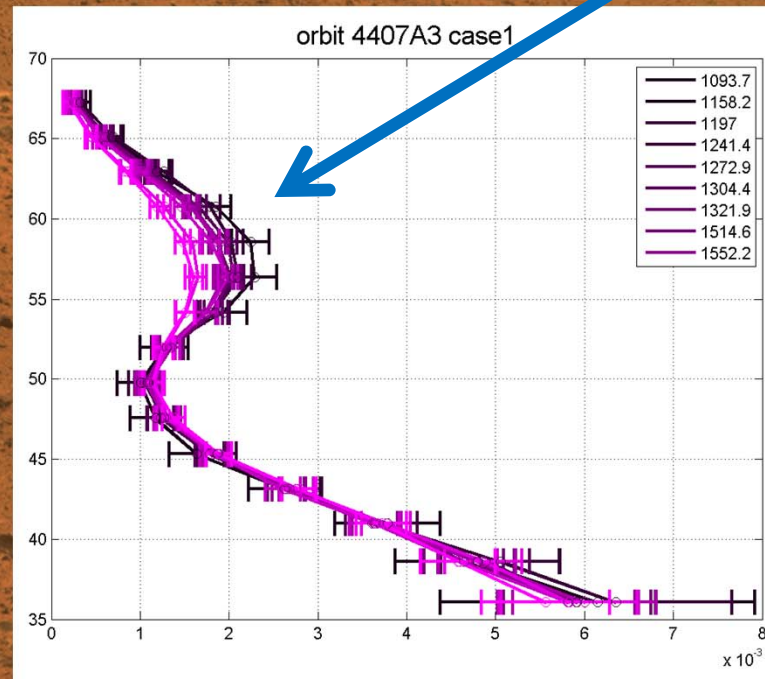
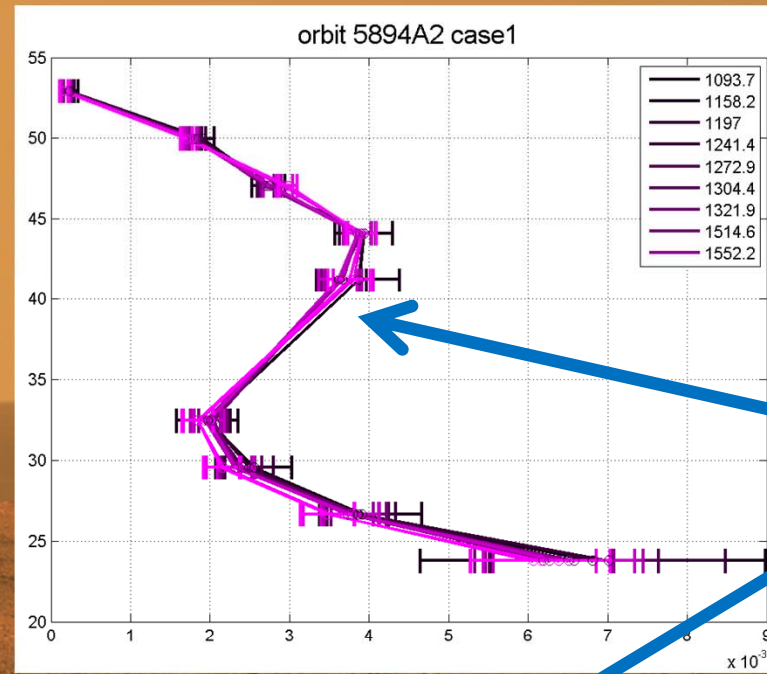
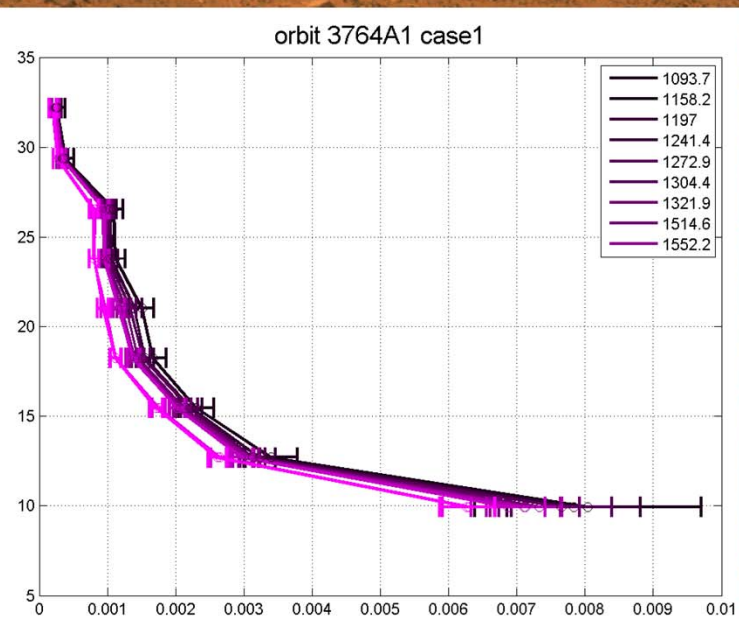
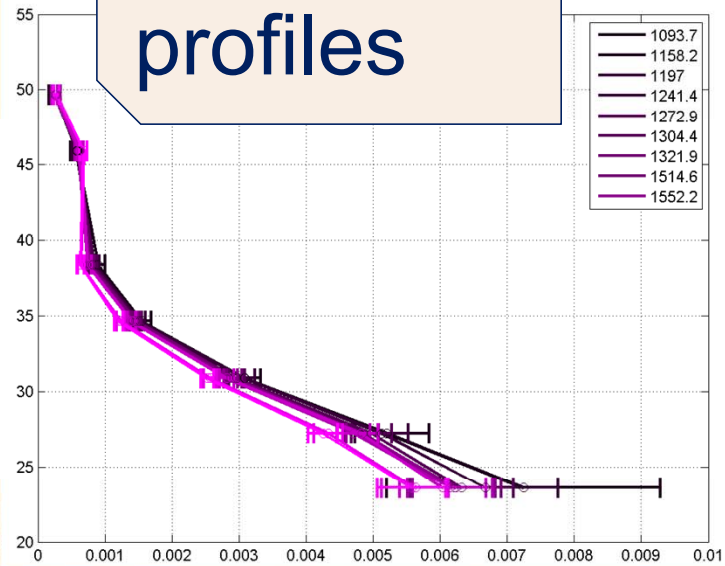
Seasonal and altitudinal dependence summarized and averaged over 7 Martian years



Where can we observe

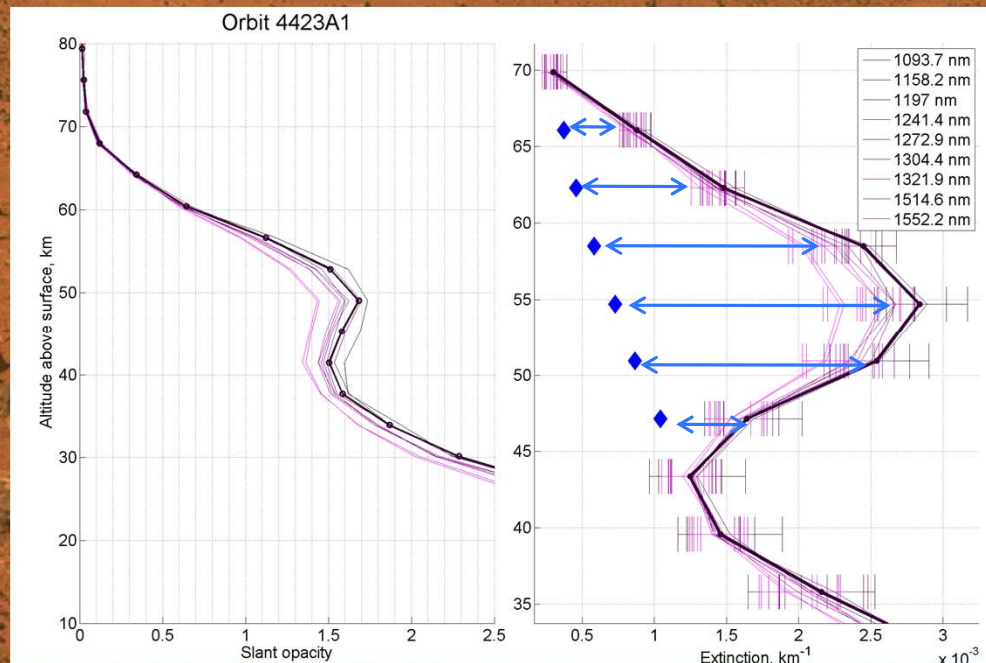
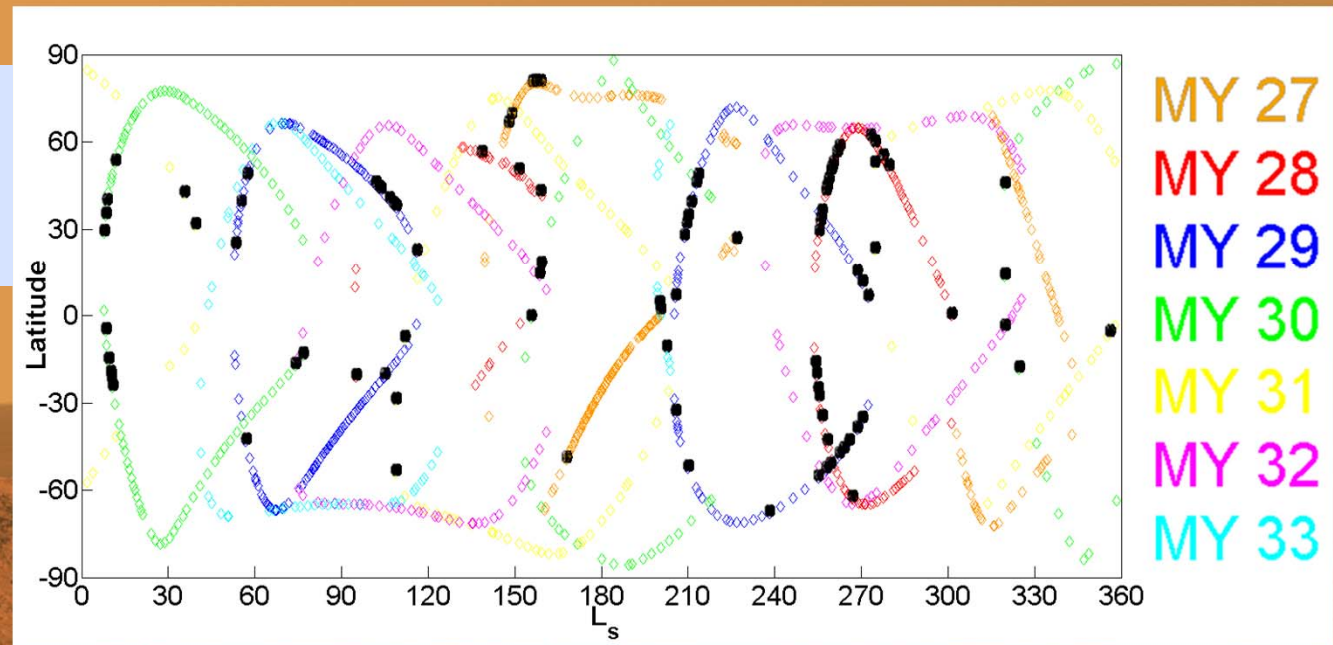


Extinction profiles



Water
ice
clouds

Water ice clouds



Particle sizes 0.8 – 1.4 μm

Opacity 0.002– 0.08

Altitudes 30 – 60 km

Mostly the **size distribution** is narrow

$$\nu_{eff} = 0.1 - 0.3$$

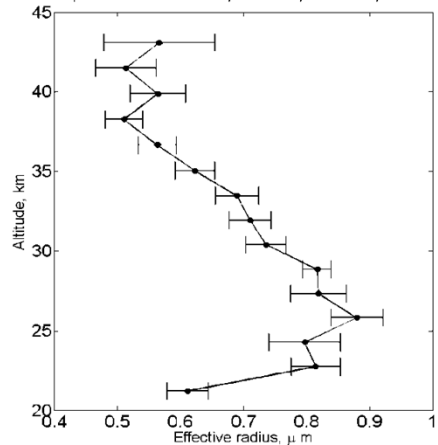
Refractive indices for water ice are taken from:

Warren, S.G., Brandt, R.E., 2008. **Optical constants of ice from the ultraviolet to the microwave: A revised compilation.** J. Geophys. Res. 113, D14220.

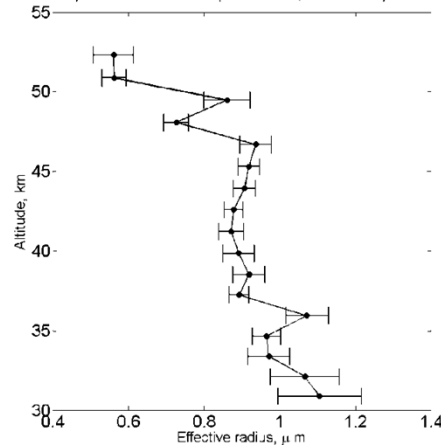
<http://dx.doi.org/10.1029/2007JD009744>

Size distribution: effective radius

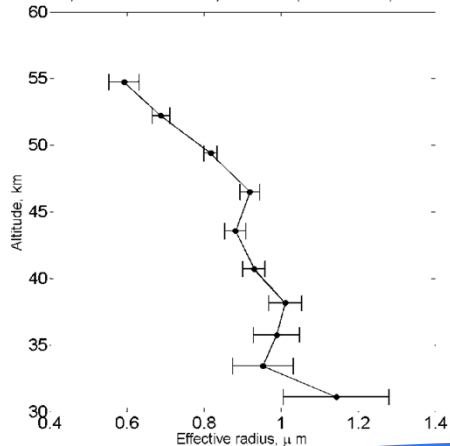
MY 29, orbit5457A2case1, Ls = 53, Lat = 20, Lon = 348



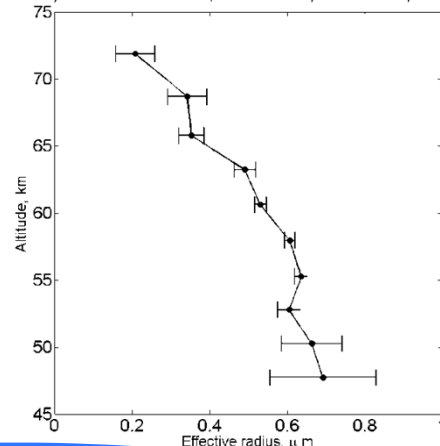
MY 30, orbit7513A1case1, Ls = 7.74, Lat = 4.21, Lon = 73



MY 30, orbit7520A1case1, Ls = 8.71, Lat = -8.56, Lon = 91



MY 33, orbit15966A1case1, Ls = 199.53, Lat = 11.42, Lon = 340



Samples of profiles at low latitudes: particles ~ 1 μm

$$n(r) = \frac{Const}{r} \cdot \exp\left(-\frac{(\ln r - \ln r_g)^2}{2 \ln^2 \sigma_g}\right)$$

$$r_{eff} = r_g \exp\left(\frac{5}{2} \ln^2 \sigma_g\right)$$

The method is described in detail in:

Fedorova, A.A., Korablev, O.I., Bertaux, J.-L., Rodin, A.V., Montmessin, F., Belyaev, D.A., Reberac, A., 2009.

Solar infrared occultations by the SPICAM experiment on Mars Express: Simultaneous observations of H₂O, CO₂ and aerosol vertical distribution. *Icarus*, 200 (1), 96–117. DOI:

<http://dx.doi.org/10.1016/j.icarus.2008.11.006>

Calculation algorithm:

Mishchenko, M.I., Dlugach, J.M., Yanovitskij, E.G., Zakharova, N.T., 1999. **Bidirectional reflectance of flat, optically thick particulate layers: And efficient radiative transfer solution and applications to snow and soil surfaces.** *Journal of Quantitative Spectroscopy and Radiative Transfer*, 63, 409–432.

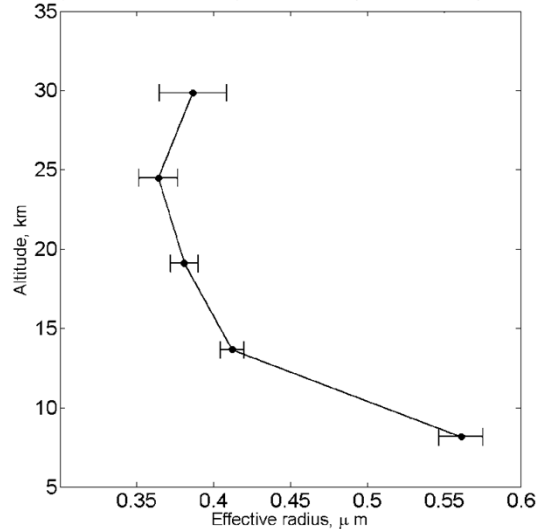
DOI: [http://dx.doi.org/10.1016/S0022-4073\(99\)00028-X](http://dx.doi.org/10.1016/S0022-4073(99)00028-X)

Refractive indices for mineral dust are taken from: Wolff, M.J. et al., 2009. **Wavelength dependence of dust aerosol single scattering albedo as observed by the Compact Reconnaissance Imaging Spectrometer.** *J. Geophys. Res.* 114, E00D04. DOI:

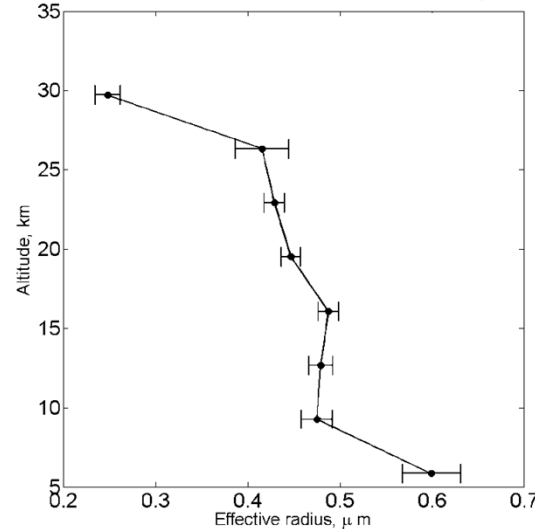
<http://dx.doi.org/10.1029/2009JE003350>

Size distribution: effective radius

MY 30, orbit8776A2case1, Ls = 180.06, Lat = -83.61, Lon = 159

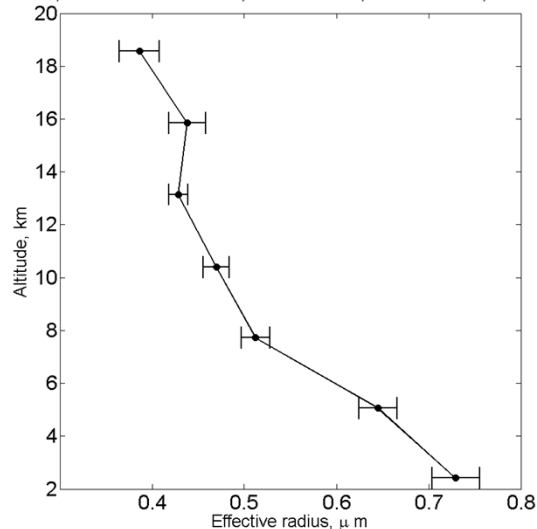


MY 30, orbit8644A3case1, Ls = 159.08, Lat = -65.34, Lon = 341

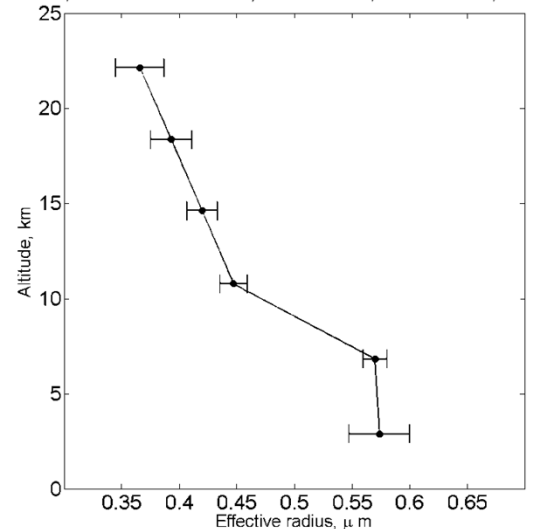


Samples of profiles at high latitudes in the Southern hemisphere during the Southern winter (**calm season**): particles $\sim 0.3\text{--}0.7\ \mu\text{m}$

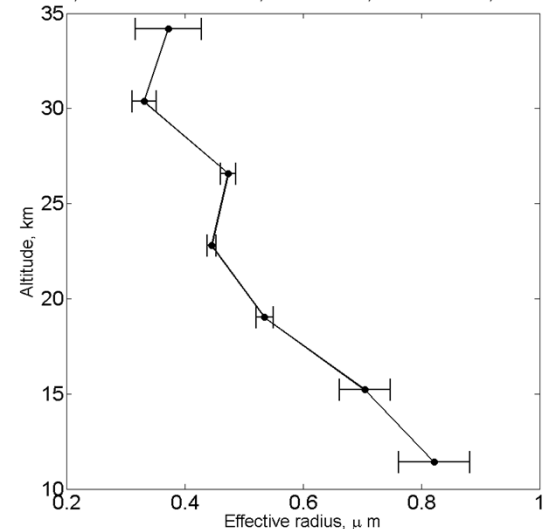
MY 31, orbit10670A2case1, Ls = 111.82, Lat = -61.90, Lon = 152



MY 31, orbit10703A2case1, Ls = 116.25, Lat = -65.45, Lon = 34



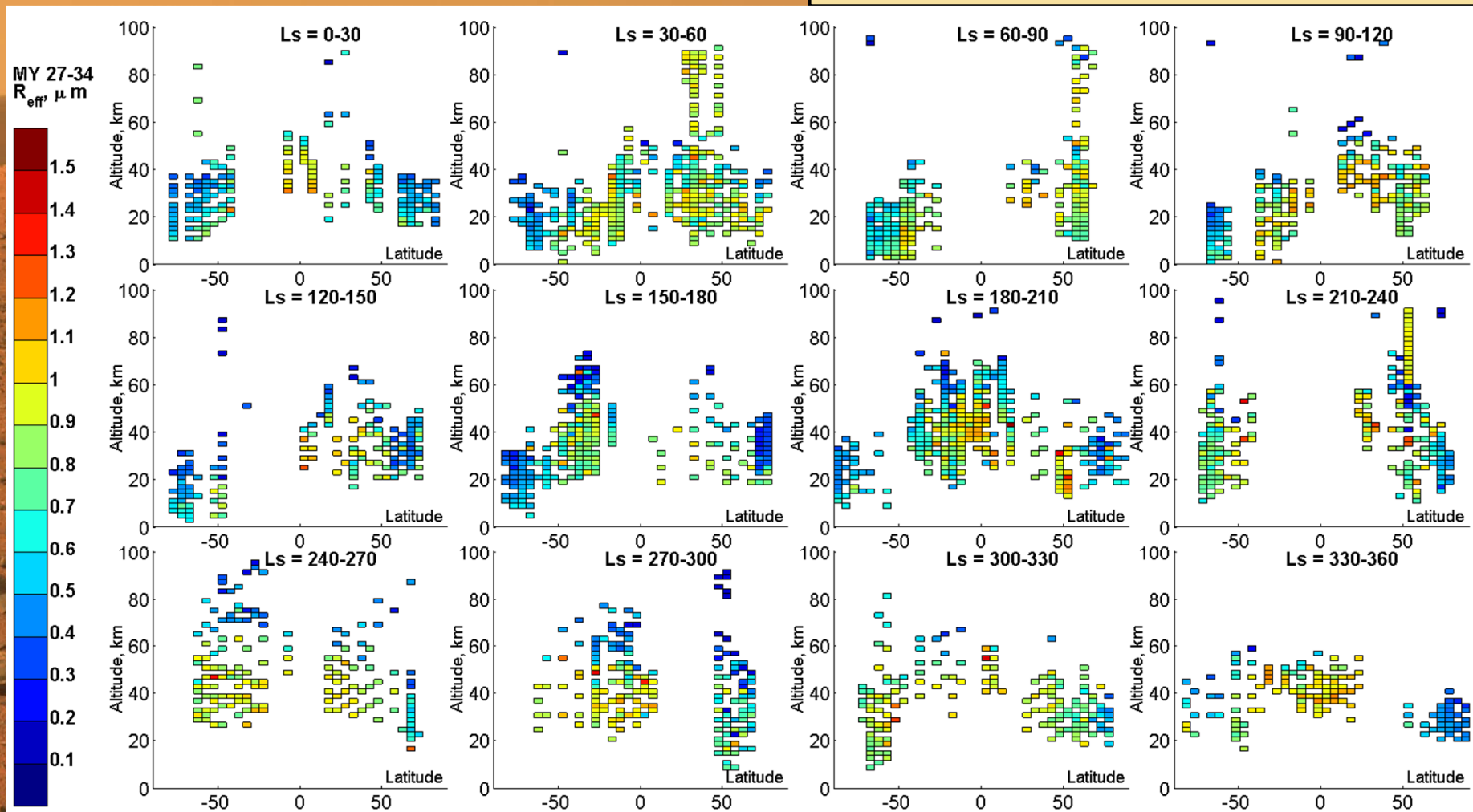
MY 34, orbit16975A2case1, Ls = 11.28, Lat = -63.09, Lon = 282



Size distribution: effective radius

$$n(r) = \frac{Const}{r} \cdot \exp\left(-\frac{(\ln r - \ln r_g)^2}{2 \ln^2 \sigma_g}\right)$$

$$r_{eff} = r_g \exp\left(\frac{5}{2} \ln^2 \sigma_g\right)$$

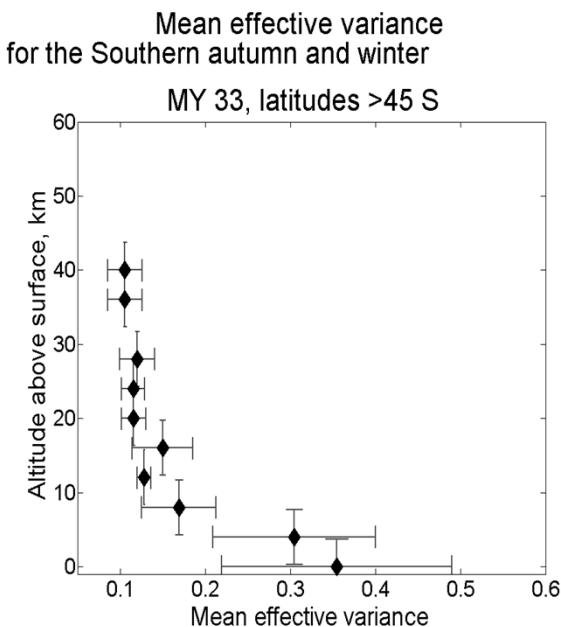


Size distribution: effective variance

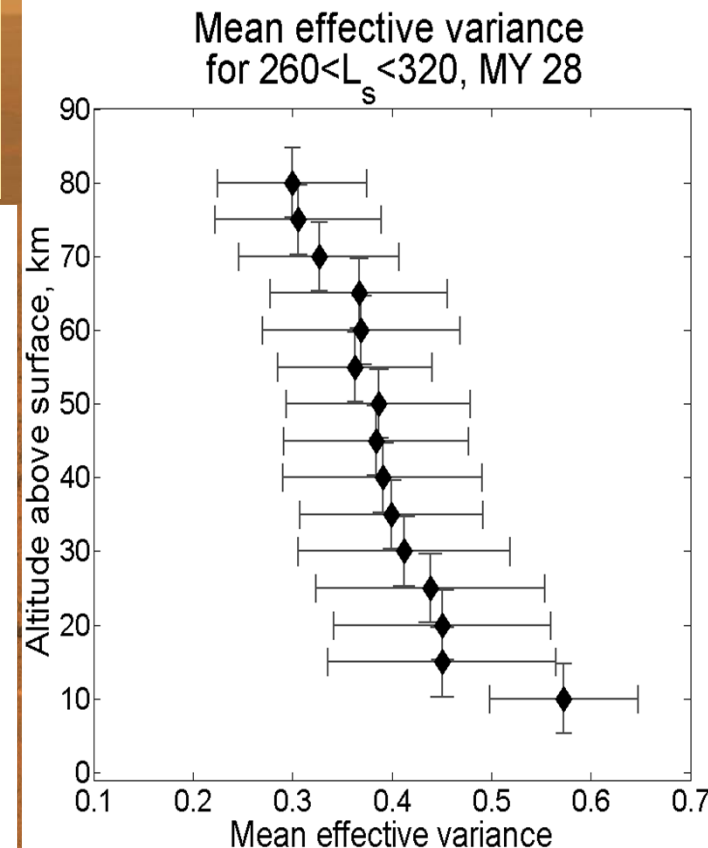
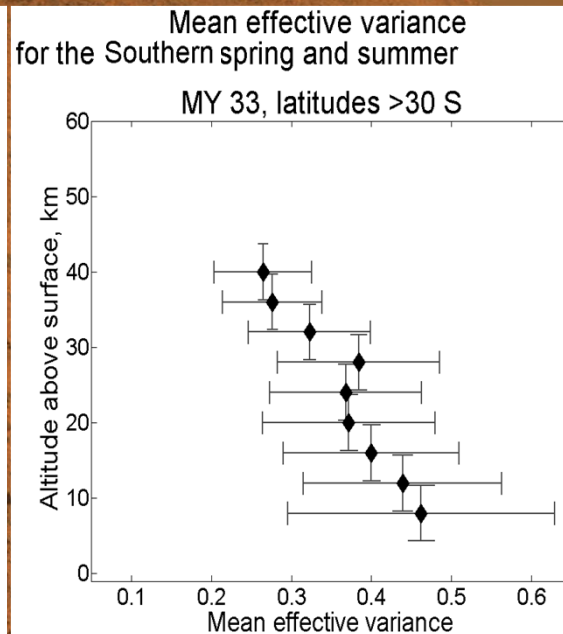
$$n(r) = \frac{Const}{r} \cdot \exp\left(-\frac{(\ln r - \ln r_g)^2}{2 \ln^2 \sigma_g}\right)$$

$$\nu_{eff} = \exp(\ln^2 \sigma_g) - 1$$

High latitudes in the Southern hemisphere,
Southern autumn and winter
(**calm season**)

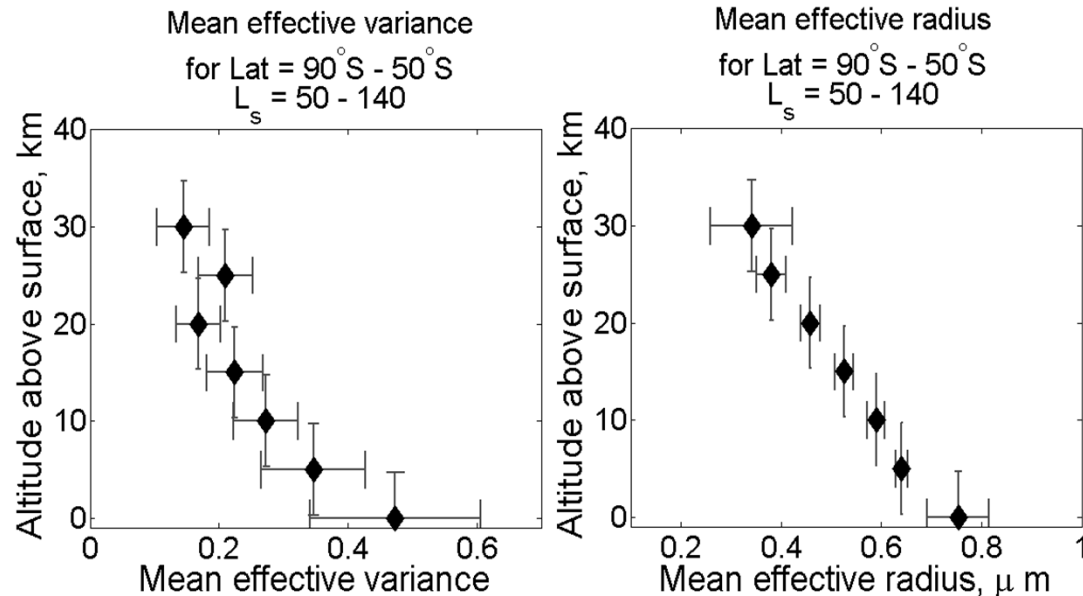


High and middle latitudes in
the Southern hemisphere,
Southern spring and summer
(**dust season**)



Global dust storm 28
Martian Year

Size distribution

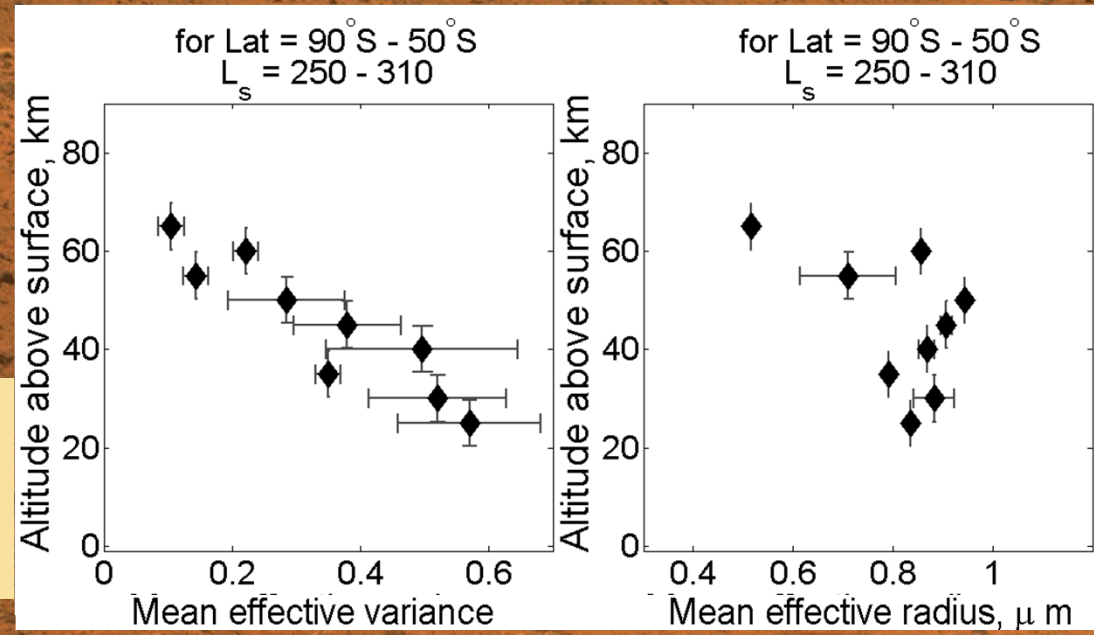


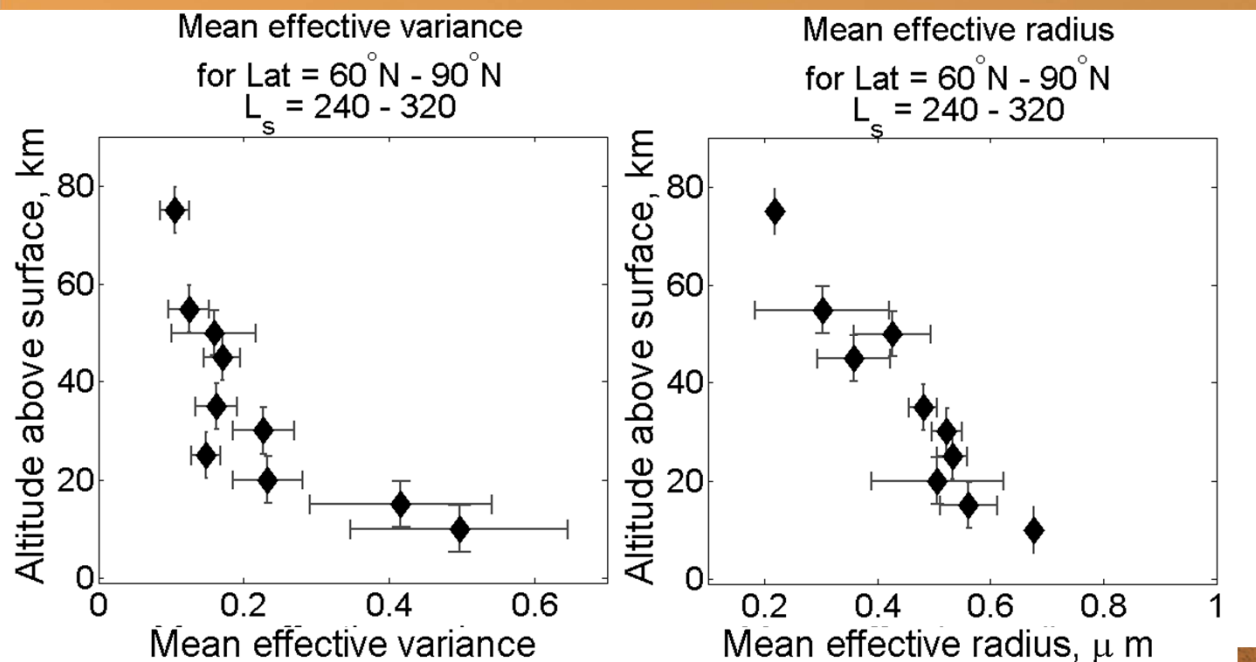
We can retrieve the averaged profiles for the different seasons: for effective radius and variation of distribution

← High latitudes in the Southern hemisphere, Southern winter (**calm season**). **Small particles, narrow distribution** even above 15-20 km



High and middle latitudes in the Southern hemisphere, Southern summer (**dust season**): **Large particles, wide distribution**



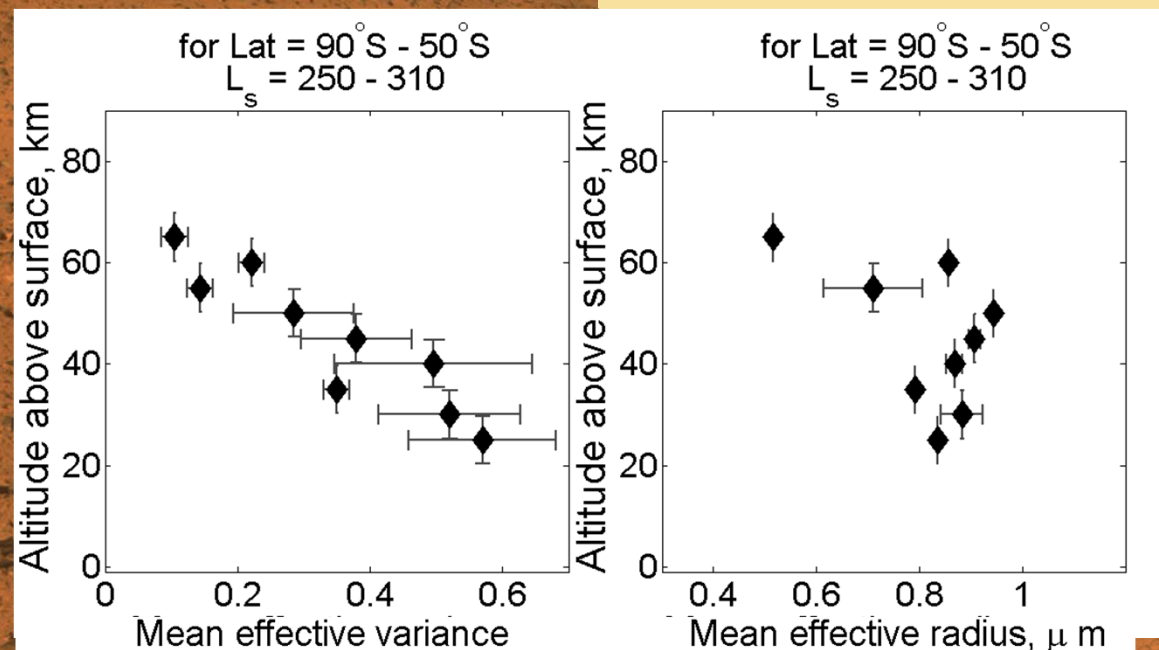


High latitudes in the Northern hemisphere,
Northern winter:
narrow distribution

High and middle latitudes in the Southern hemisphere,
Southern summer (**dust season**):
large particles, wide distribution

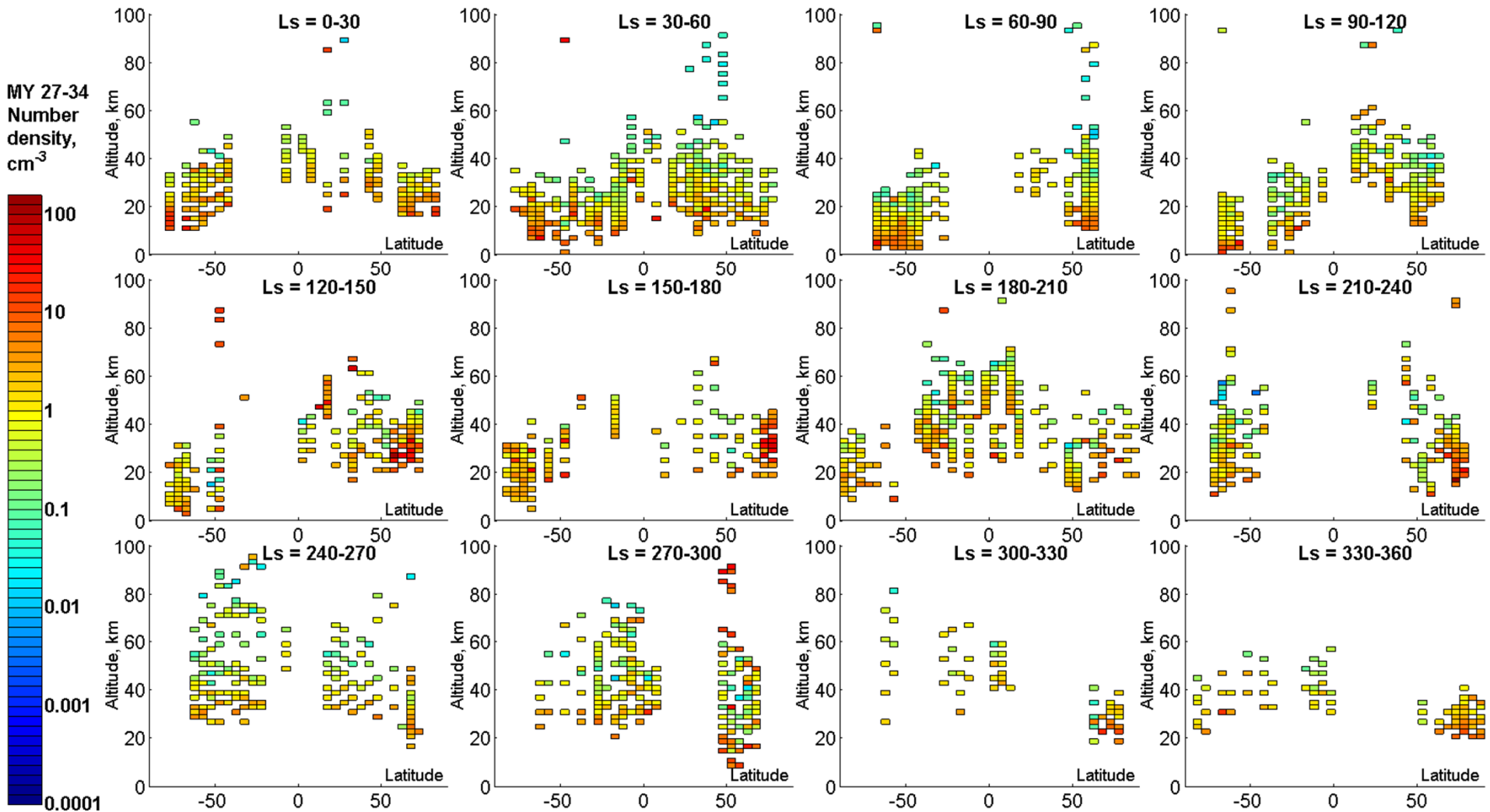


Size distribution



Number density

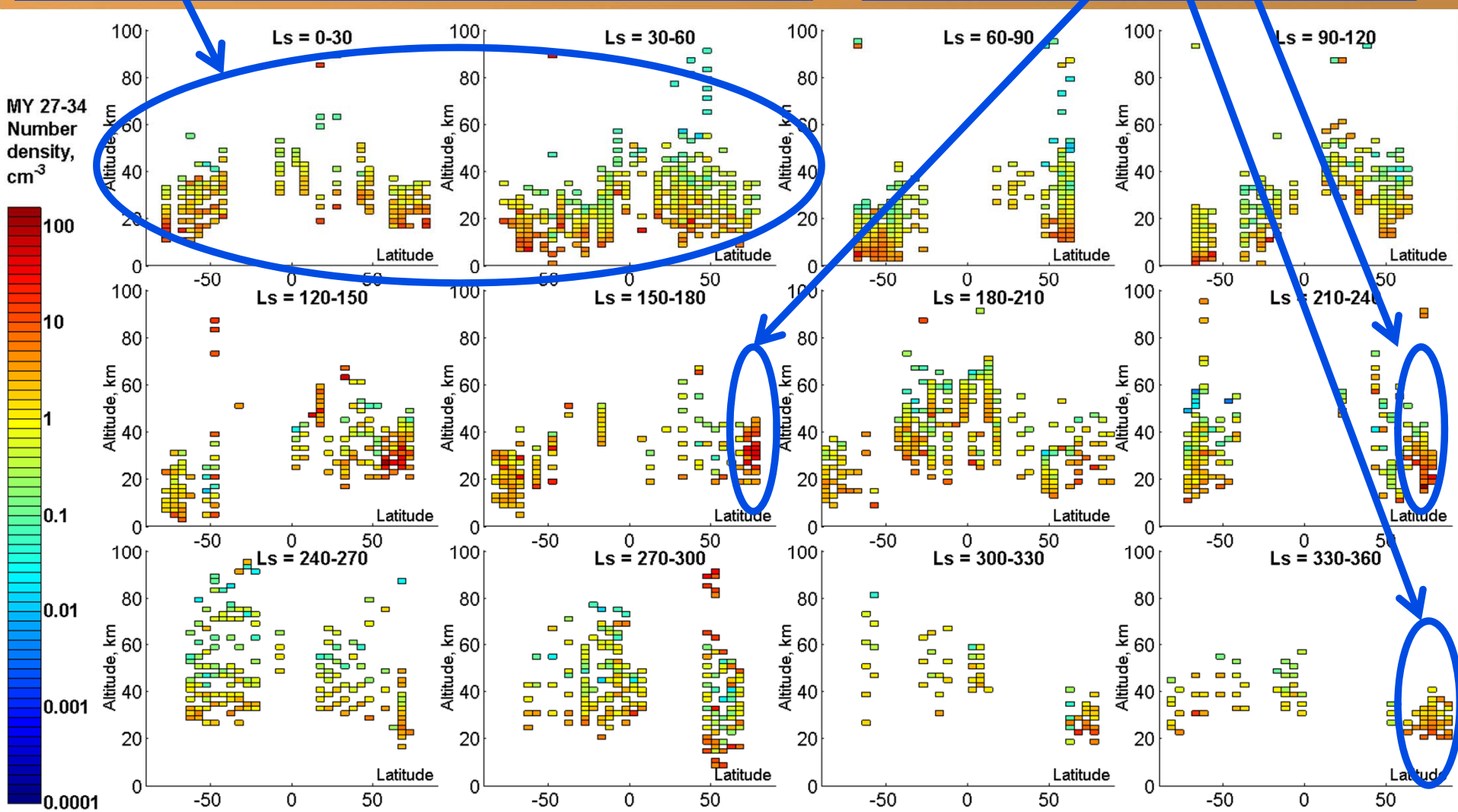
$$N(z) = \frac{k_{ext}(\lambda, z)}{\int_{r_1}^{r_2} \pi r^2 \tilde{Q}_{ext}(r, m, \lambda) n(r, z) dr}$$



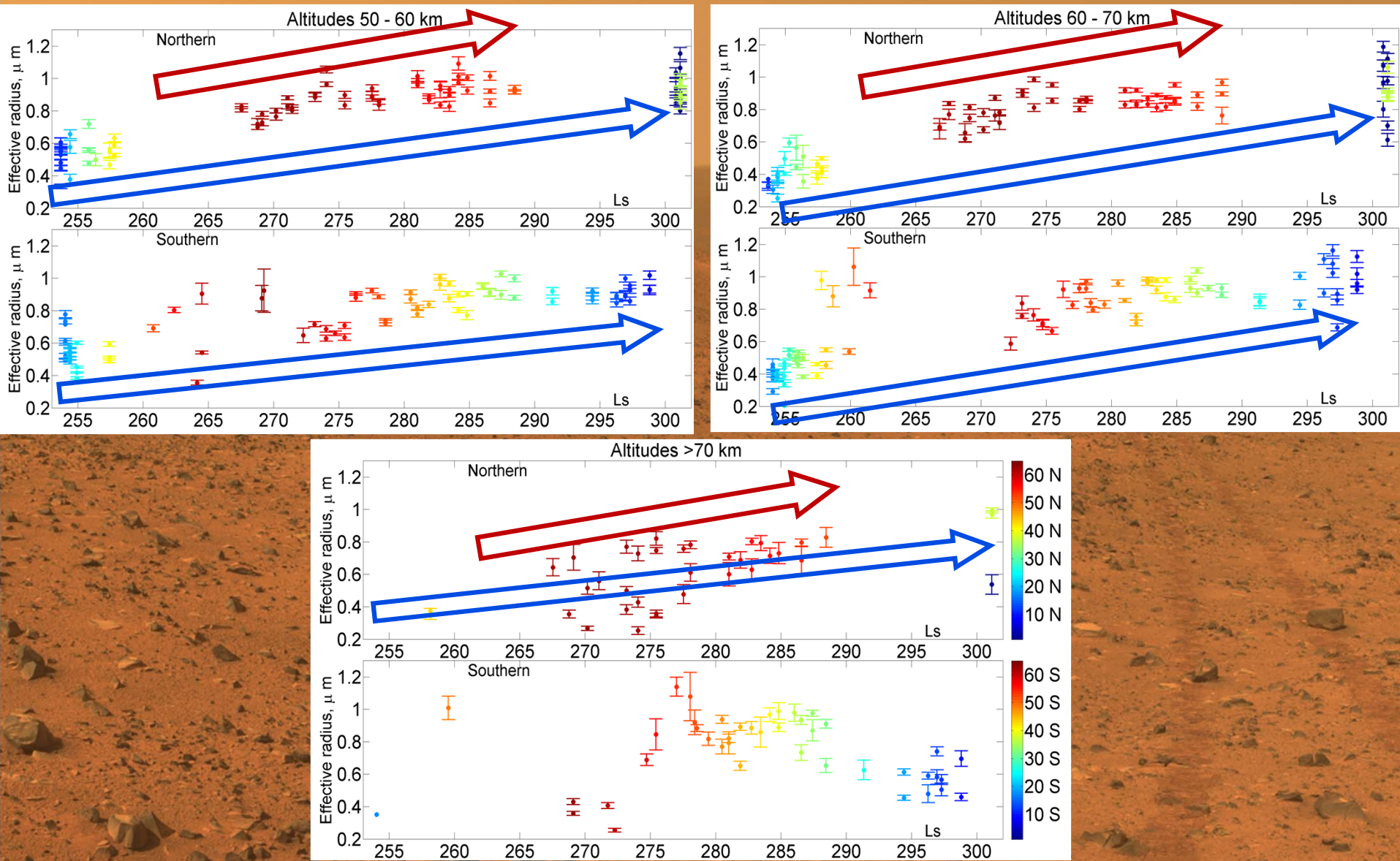
Number density

There is no clear dependence on latitude;
Number density becomes higher close to surface

High concentration of small particles at high latitudes



Global dust storm - 28 MY

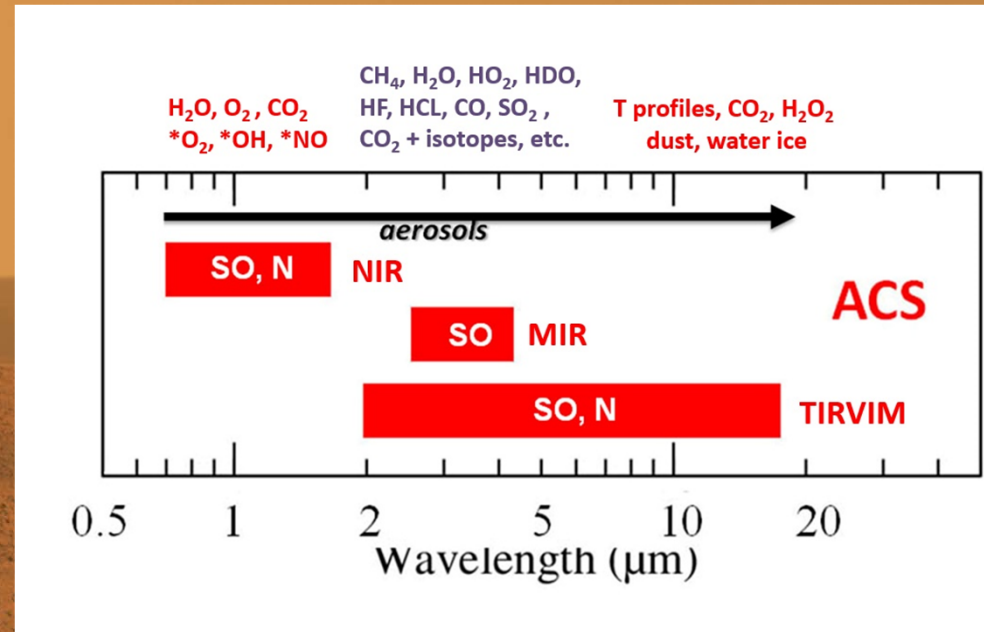


From Mars Express to ExoMars

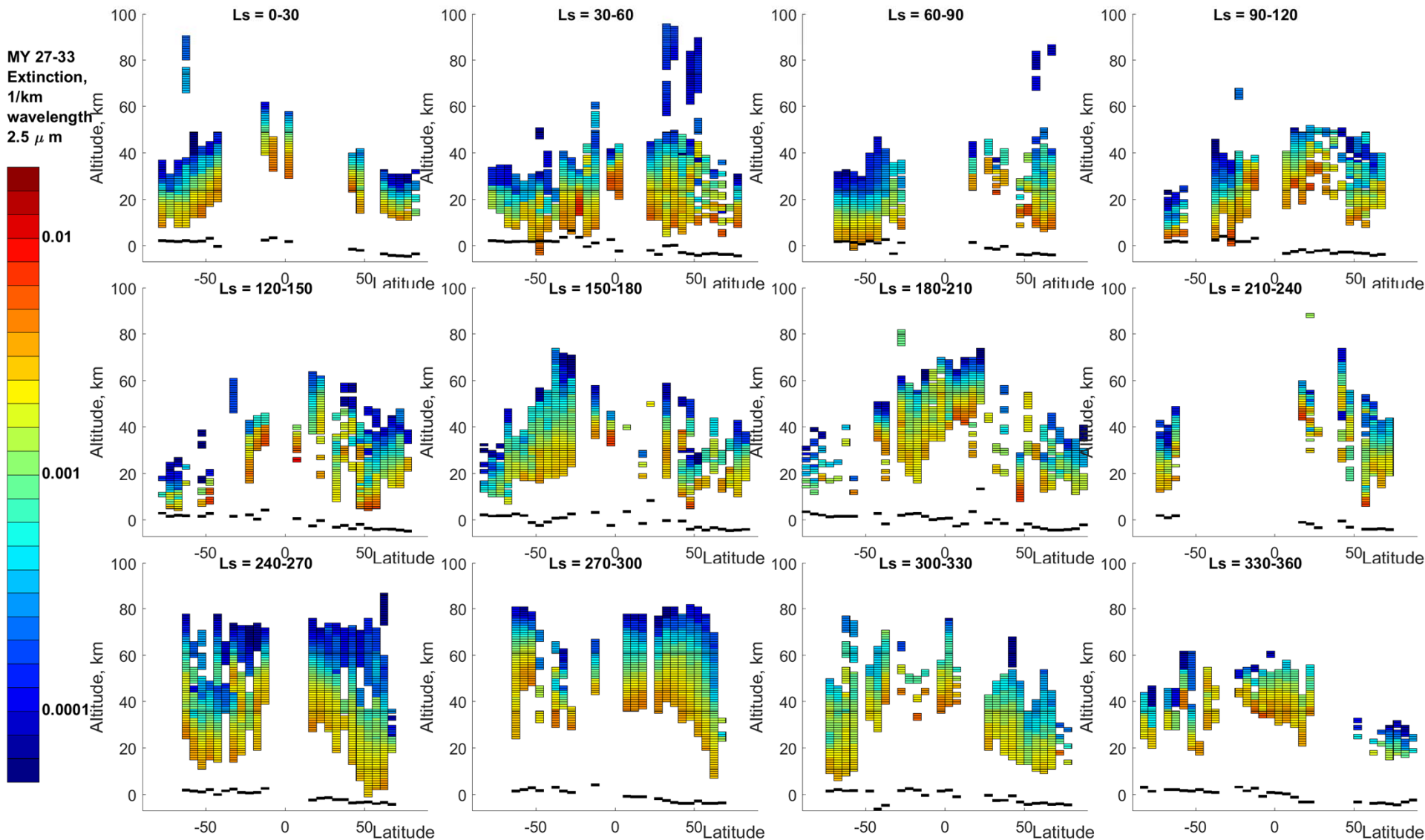
ACS - Atmospheric Chemistry Suite, three infrared instruments to investigate the chemistry and structure of the Martian atmosphere.

→ Using the retrieved size distribution, refractive indices and Mie theory we can calculate the extinction coefficient as:

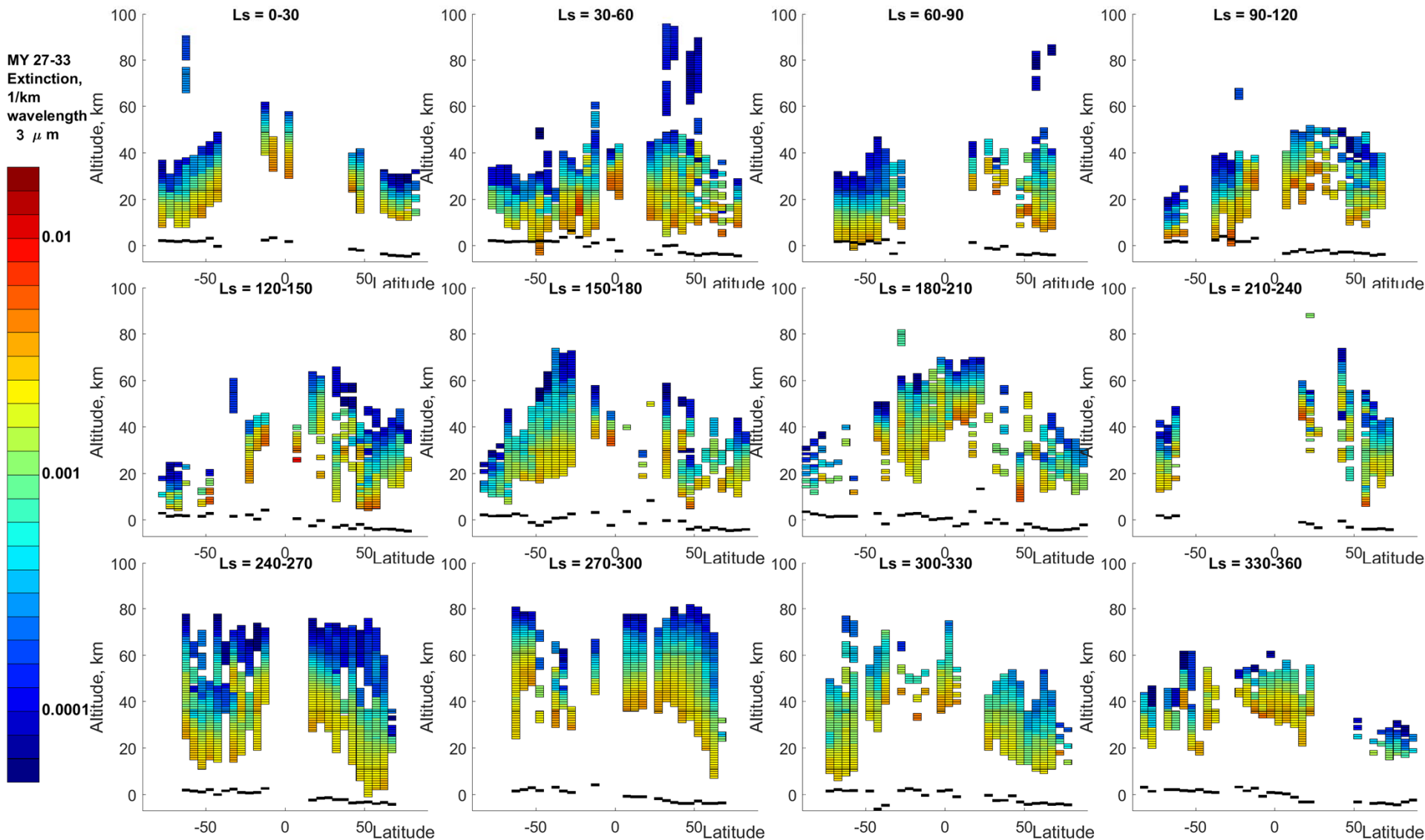
$$k_{\text{ext}}(\lambda, z) = \int_0^{\infty} \sigma_{\text{ext}}(r, z) n(r, z) dr$$



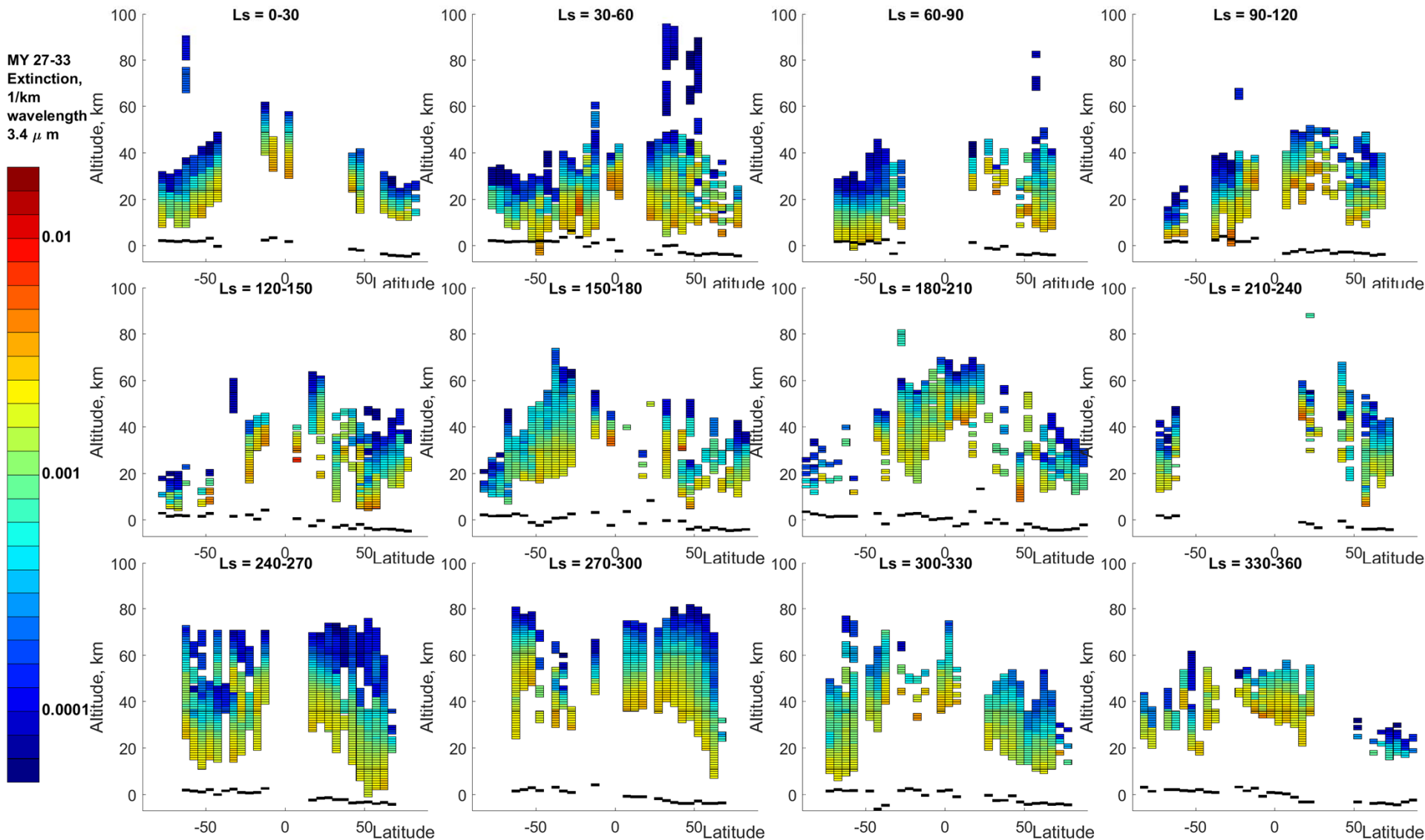
Extinction coefficient map on 2.5 μm (calculated using retrieved particle size distribution)



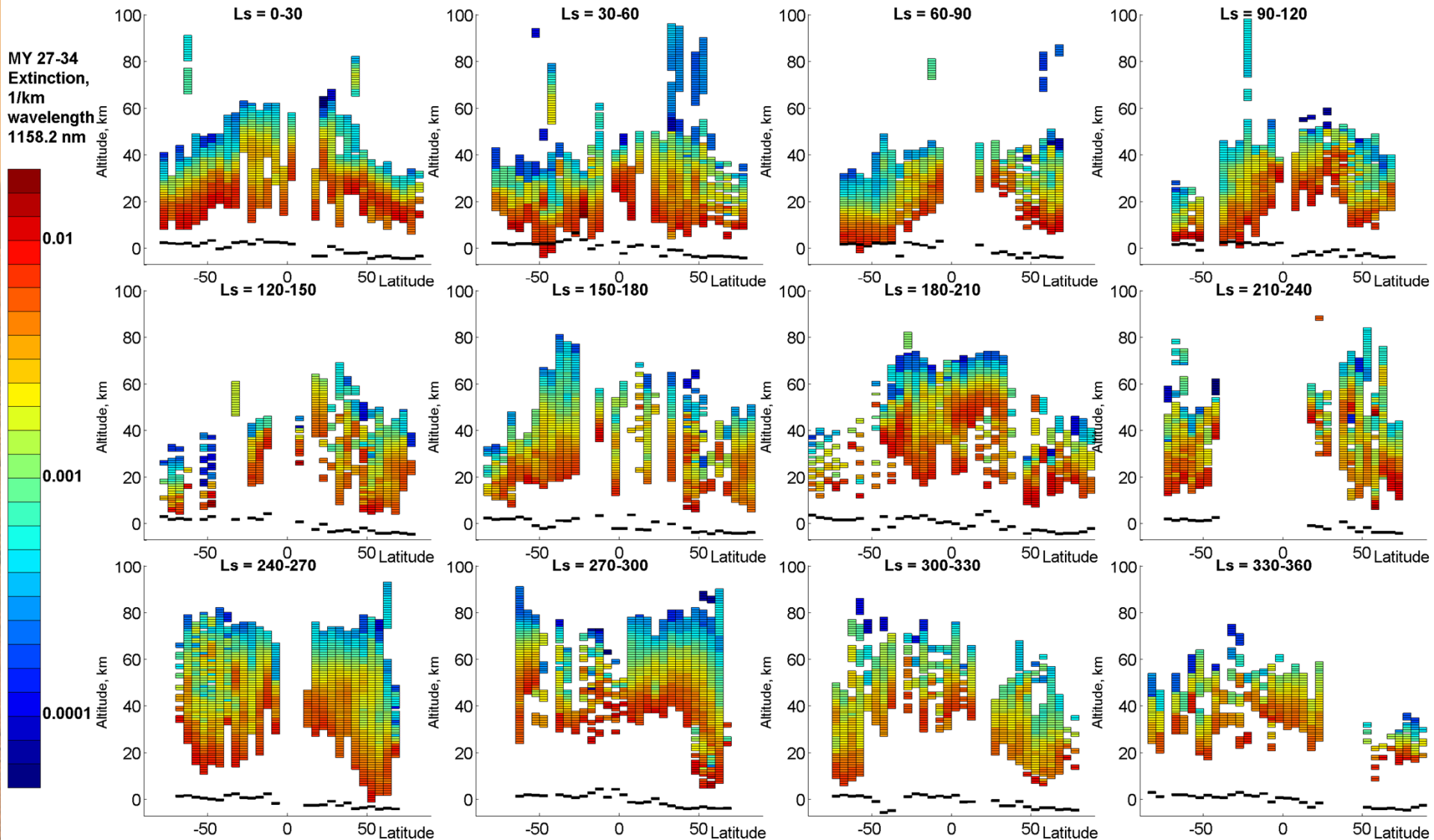
Extinction coefficient map on 3 μm (calculated using retrieved particle size distribution)



Extinction coefficient map on 3.4 μm (calculated using retrieved particle size distribution)



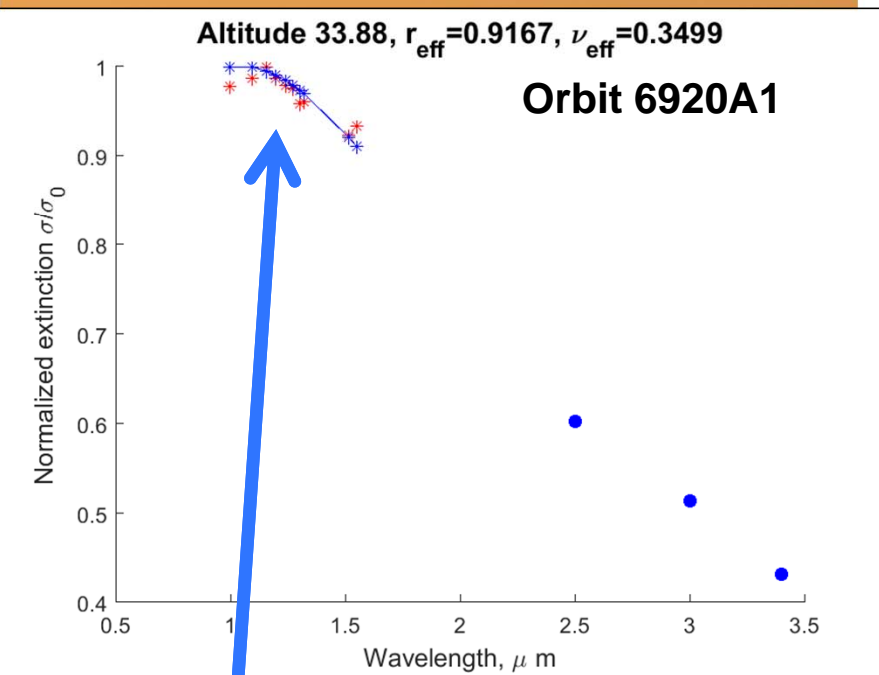
Extinction coefficient map on 1.158 μm





Thank you for your attention!

Calculation for extinction on 2.5 μm , 3 μm and 3.4 μm



Fitting SPICAM IR data

