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Thermospheres of Terrestrial Planets Including Coupling with the Lower Atmosphere

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Outline of Talk

Basic Thermospheric Features and P. ParametersThree Common Features and Processes to Contrast

- * Venus and Mars NO nightglow distributions as a tracer of UATM winds (proxy view of global circulation patterns).
- * Wind induced diffusion and Helium distributions
- * Day thermal budgets and relative roles of CO_2 15-µm cooling, molecular thermal conduction and global winds

Lower to Upper Atmosphere Coupling

- * Migrating and non-migrating tides of Mars vs Earth (densities, nightglow)
- * Homopause altitudes and eddy diffusion K_{zz} for Mars vs Earth
- * Dust storms on Mars (regional storm impacts)
- * In-situ Mars thermospheric wind measurements (non-solar forcing)
- * Gravity waves at Venus (see talk by A. Brecht)



Basics Thermospheric Features and Planetary Parameters



Fundamental Planetary Parameters: Some Implications

Parameter	Venus	Earth	Mars
Gravity (cm ² /sec): H(km)	888 [4-12]	982 [10-50]	373 [8-22]
Heliocentric distance (AU)	0.72	1.0	1.38-1-67
Radius (km)	6050	6371	3390
Omega (rad/sec): Coriolis	3.0(-7) [no]	7.3(-5) [yes]	7.1(-5) [yes]
Obliquity (deg): Seasons	1-3 [no]	23.5 [yes]	25.12 [yes]
B-dipole moment (wrt Earth): Auroral/Joule heat	≤4.0(-5) [no]	1.0 [yes]	≤2.5(-5) [no] Crustal fields
EUV-UV (eff.)	~20%	30-60%	~20%



Mars (and Venus) Upper Atmospheric Regions and Processes (Bougher et al., 2015)





Solar Spectral Irradiance Variations over the Solar Cycle (~0.1-190.0 nm)



Integrated 0-190 nm FISM irradiance (1947-2011)

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Solar EUV Monitor at Mars: EUVM Instrument 3-Diodes and FISM-M Daily Averaged Fluxes (Pilinski et al 2018)





Composition Differences : Venus Upper Atmosphere



Venus

PVO/ONMS Kasprzak et al (1997) Venus II book



Composition Differences : Mars Upper Atmosphere



Mars

MAVEN/NGIMS P1064, Ls ~ 256, LT ~ 12PM, LAT = 4.5S (Mahaffy et al. 2015. GRL Special Issue)

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Thermospheric Temperature Profiles (1): Venus (Pioneer Venus and Magellan)



Bougher et al. (2002)



Thermospheric Temperature Profiles (2): Earth



Adapted from Roble (1986)



Thermospheric Temperature Profiles (3): Mars (MAVEN/NGIMS Deep Dips# 1-8)



Stone et al. (2018) JGR



Three Common Features and Processes to Contrast



Statistical Maps of VEx O₂ and NO Nightglow Distributions: ~100 km & 115 km layers

Soret et al., (2012); Stiepen et al. (2013); Gerard et al. (2018)

 $O_2 \, {}^1\!\Delta_g$ statistical map (VIRTIS)

NO (gamma + delta) statistical map (SPICAV)



 $LT \sim 0 AM$

 $LT \sim 2 \text{ to } 4 \text{ AM}$

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Venus Helium Bulge at 160 km: Pioneer Venus/ONMS (Mengel et al. 1989)



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Venus Thermospheric Circulation Paradigm: Schubert et al. (2007)





VTGCM NO Nightglow VER Map (VEx/SMIN): Bougher et al. (2013); Gerard et al. (2018)



Volume Emission Rate: log_{10} (photons/cm³/s)

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MAVEN/IUVS Disk Apoapse Images of Brightness of NO Nightglow: Stiepen et al. (2018)

NO brightness disk images LS/LAT



Seasonal distribution shows brightening at winter poles with strong solstice asymmetries. Data cover full MY, with gaps. ¹⁸



Mars Helium Bulge : MAVEN/NGIMS Elrod et al. (2017)





M-GITM T+(U,V) and Helium density at ~200 km: DD2 Conditions (Ls ~ 328; SMED)

T+(U,V)

He-Density



- Low SZA temperatures peak (Texo ~ 260-270 K) along with CO_2 densities.
- High SZA (winter polar night) temperatures are a minimum along with CO₂ densities.
- Divergence (upwelling, day) and convergence (downwelling, night) of the circulation
- Winter polar night (LT ~ 2-5) peak in He densities seen, corresponding to wind transport (horizontal & vertical). He shows opposite behavior of [CO₂] distribution.



LMD-MGCM VER Map of NO Nightglow (Gonzalez-Galindo, 2018). In preparation

Ls ~ 0-30; LT = 4 (Equinox)



Ls = 270-300; LT = 4 (Perihelion)

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Units: \log_{10} (photons/cm³/s)



Diagnostic Thermal Balances using GCMs: VTGCM, TIE-GCM, and M-GITM

	VTGCM	TIE-GCM	M-GITM*
Altitude Range	~70-250 km	~95-800 km	~0-250 km
Vert. Resolution (pressure coor.)	H/2 or H/4	H/2 or H/4	H/4 typical
Horz. Resolution	5x5 or 2.5x2.5°	5x5 or 2.5x2.5°	5.0 x 5.0 $^{\circ}$ thusfar
Major Neutrals	T, U, V, W, CO ₂ , CO, N ₂ ,O	T, U, V, W, O, N ₂ , O ₂	T, U, V, W, CO ₂ , CO, N ₂ ,O
Major Ions	O ₂ +, CO ₂ +, O+, N ₂ +, NO+	O ₂ +,N ₂ +,O+, NO+, N+	O ₂ +, CO ₂ +, O+, N ₂ +, NO+
Minor Neutrals	N(⁴ S), N(² D), NO, Ar, O ₂	N(4S), N(2D), NO	N(⁴ S), Ar, He, O ₂
Rad. Cooling	CO ₂ (15-µm)	CO ₂ (15-μm) NO (5.3-μm) O (63-μm)	CO ₂ 15-µm О (63-µm)



Venus Dayside-Nightside Exospheric Temperatures: Kasprzak et al. (1997) Pioneer Venus



 $\Delta T \sim 70-80 \text{ K}$ (at Venus)

Nearly Constant

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VTGCM T-Profiles: Dayside and Nightside for SMIN and SMAX Conditions





VTGCM Thermal Balances at SZA = 0: (SMIN/VEX and SMAX/PVO Conditions)





Mars Texo Variations Over Solar Cycle: Bougher et al (2009) Pre-MAVEN





IUVS Dayglow Scale Height (~150-180 km) vs Lyman-α Irradiance (1-MY): Jain et al. (2018)



- Temperatures (scale heights) are highly correlated with incoming solar EUV
- Scale heights are also highly correlated with Sun-Mars distance

DD2 versus DD8 Averaged Dayside T- Profiles: NGIMS-MGITM Comparisons + Thermal Balances



Texo ~ 270 K (MGITM) and ~265 K (Stone et al. 2018): DD2 Texo ~ 190-200 K (MGITM) and ~190-200 K (Stone et al. 2018): DD8 ΔTexo ~ 65-70 K (DD2: SMED/EQU to DD8: SMIN/APH)



Lower to Upper Atmosphere Coupling with Examples from Mars Spacecraft Observations



Summary of Primary Tidal Forcings for Earth and Mars (Siskind & Bougher, 2016)

Tidal Component	Earth	Mars
Migrating diurnal	NIR abs. of H_2O	VIS and IR abs of CO_2
Migrating semi-diurnal	Ozone heating	Dust heating
Non-migrating	Latent heat release from tropical convection	Topographical and pl. wave modulation of near surface heating

- Similar rotation rates of Earth and Mars imply that the mathematical characteristics of the tidal modes for Earth and Mars are similar.
- In general, due to large topographic variations, the manifestation of these tides in the Martian upper atmosphere is greater on Mars than Earth.
- Overall: Local time sorting is needed to go beyond harmonic analysis to address identification of specific tides.



MAVEN IUVS CO₂⁺ UVD Derived CO₂ Density Variations in Longitude (Lo et al. 2015)



Harmonic fits to LON variations in CO_2 densities as observed by IUVS at 130 km. Wave-2 has largest amplitude (~20-30%) and near constant phase at low LAT.

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IUVS CO₂⁺ VER and NGIMS CO₂ Variations in Longitude (England et al. 2016) IUVS NGIMS



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IUVS NO Nightglow Periapse Data & Disk Images Show Wave-3 Equatorial Structure during Southern Winter (Royer et al. (2017)

Periapse data: 80 < Ls < 120



Disk images: 130 < Ls < 170



~1 Martian Year apart Same equatorial latitude range

We observe:

Changes in the longitudes of the peaks,

Changes in the space between each peak,

Changes in the ratio between enhanced regions



Mars Homopause Altitudes Derived from NGIMS N₂/Ar Ratios: Slipski et al. (2017)



- Homopause varies significantly and does not track a fixed density level
- Lowest homopause altitudes occur between 4 6 AM and the highest values near noon when Ls > 270° (near perihelion). Implies Kzz ~ 10 to 10,000 m²/s



Dust Storm Impacts on Mars Thermosphere: Regional Storms

MY23 Noachis Storm (Ls ~ 224): Mass Density at 130 km (Bougher et al., 1999) MY27 Storm (Ls ~ 130): Ion Peak Height (Withers & Pratt, 2013)



MGS Orbit Number

MGS Radio Occultations

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**NGIMS sampling during regional storms (Jan-Apr 2017) and impacts over 170-220 km (Liu et al., 2018). ~50-200% density enhancements.



Thermospheric Neutral Winds Derived From NGIMS Observations (Roeten et al., 2018)

250.0 m/s



- Winds were derived below ~200 km using technique of "yawing" NGIMS left/right; monthly 2-day (10-orbit) campaigns since September 2016
- Large, unexpected orbit-to-orbit measured wind variability, as large as ± 100 m/s
- Comparisons with M-GITM model predictions show good agreement much of the time. However, significant differences suggest the importance of processes such as vertically propagating gravity waves and tides. Cite examples above.



Major Take Aways

- □ Planetary fundamental parameters indeed play a significant role in determining UATM structure and dynamics on these terrestrial planets. Exoplanet (HZ) simulations possible using these planets as prototypes.
- □ Solar cycle variation of dayside Texo smallest (~70-80K) for Venus, owing to thermostatic control by CO₂ cooling. Mars day Texo are instead regulated by large scale dynamics, CO₂ cooling and thermal conduction, yielding larger variations (~70-140K) over the solar cycle.
- □ Coupling of 3-terrestrial planet thermospheres (from below) is very important for determining structure and dynamics (e.g. tides & dust).
- □ NO nightglow is being used effectively at Venus and Mars to serve as a tracer of UATM circulation and a monitor of short term variability. NGIMS wind measurements suggest coupling from below is likely.
- □ Mars homopause altitudes show large variations with season, latitude and local time. Consistent with large K_{zz} (up to 10⁸ cm²/s). 100x Earth.

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