



Metastable oxygen O(¹S) Martian airglow: observations and model

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Why study this particular emission?

- Not particularly bright
- Not extensively observed in the past

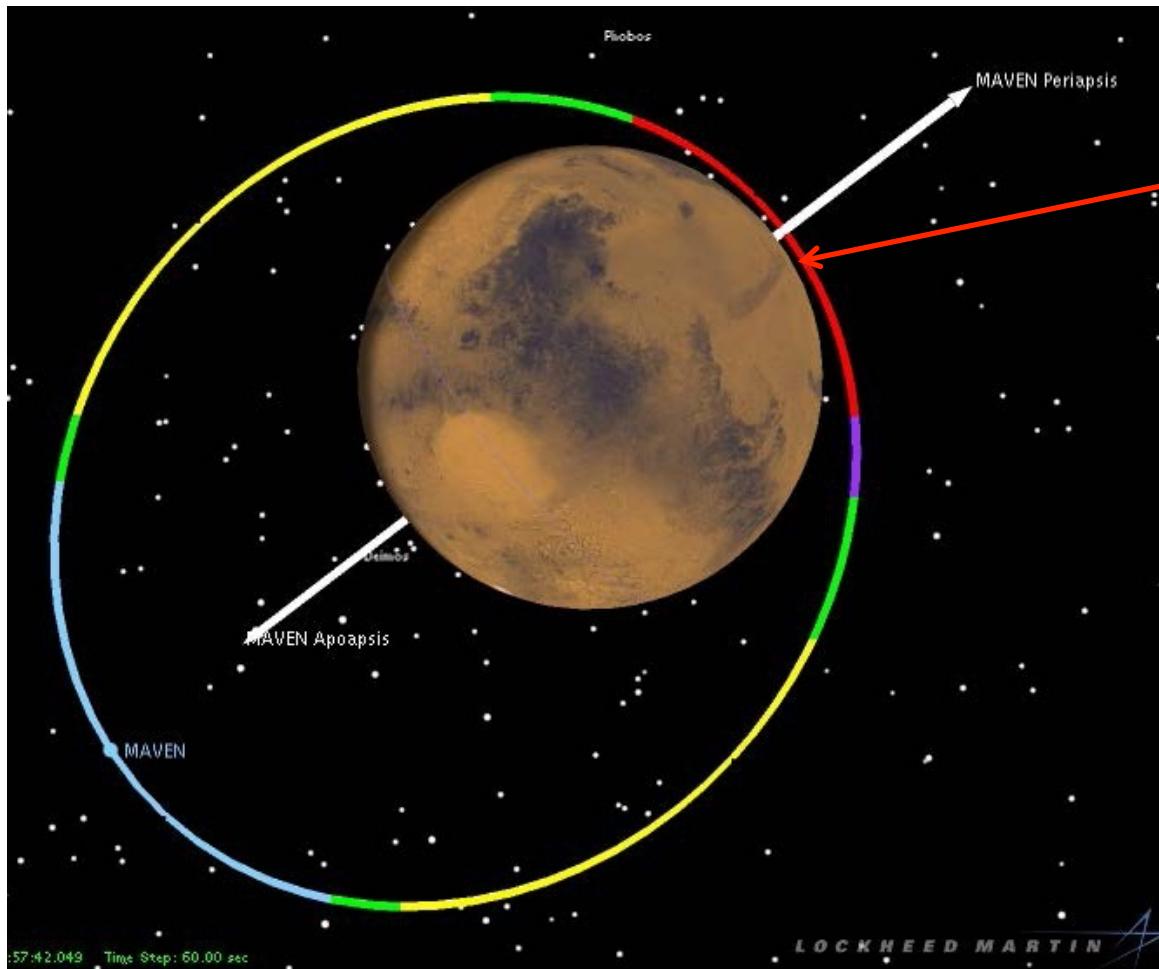
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BUT:

- IUVS-MAVEN improved sensitivity provides high S/N ratio
- Its intensity distribution is relatively simple to model and provides direct information on the thermospheric structure
- Very close behavior to a Chapman layer

MAVEN orbit

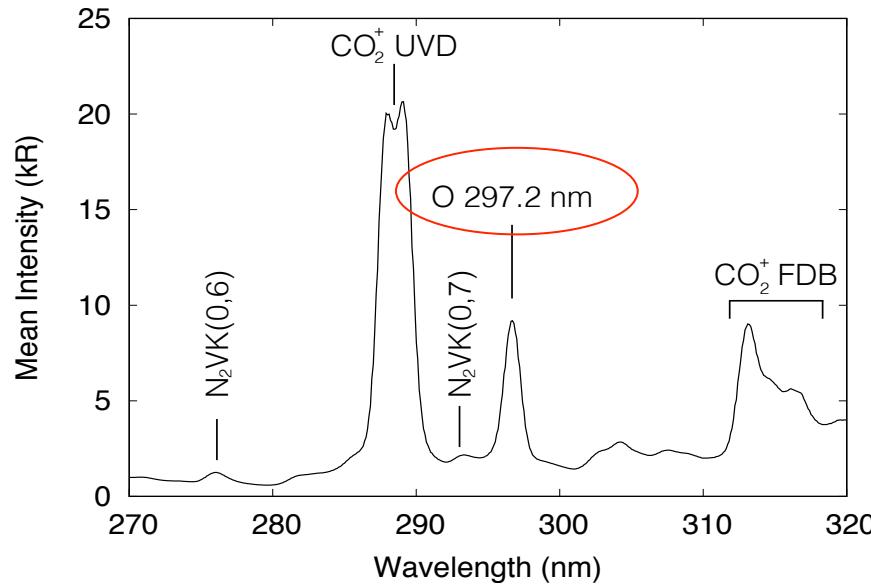


Periapsis phase:
up to 12 successive
limb scans

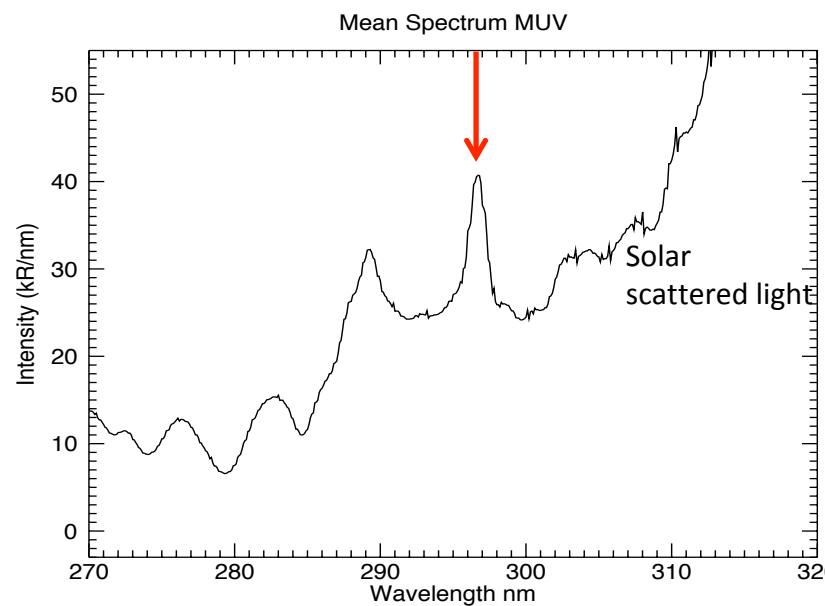
IUVS spectral range:
120 to 320 nm

IUVS Resolution:
1.2 nm

Sum of IUVS limb spectra in the region of the O 297.2 nm emission

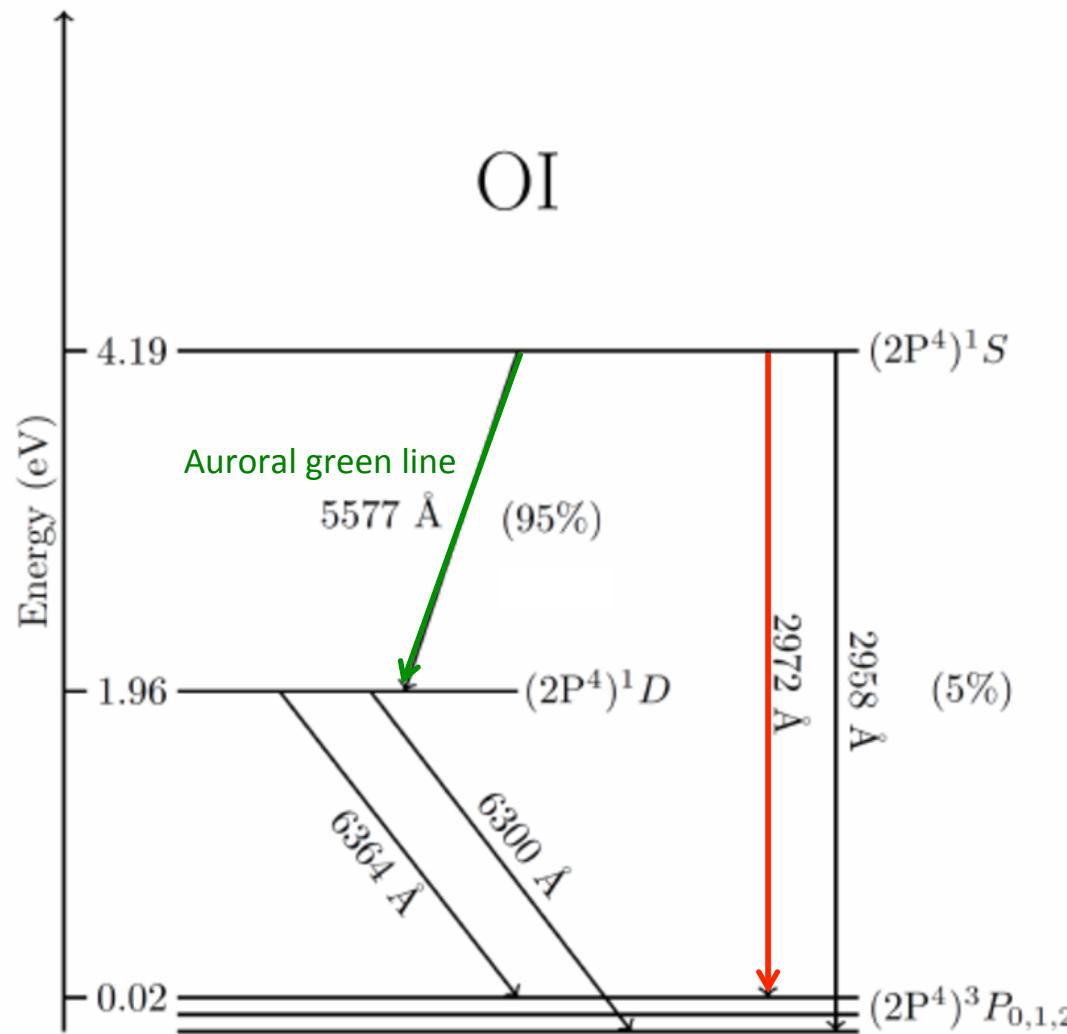


130 km



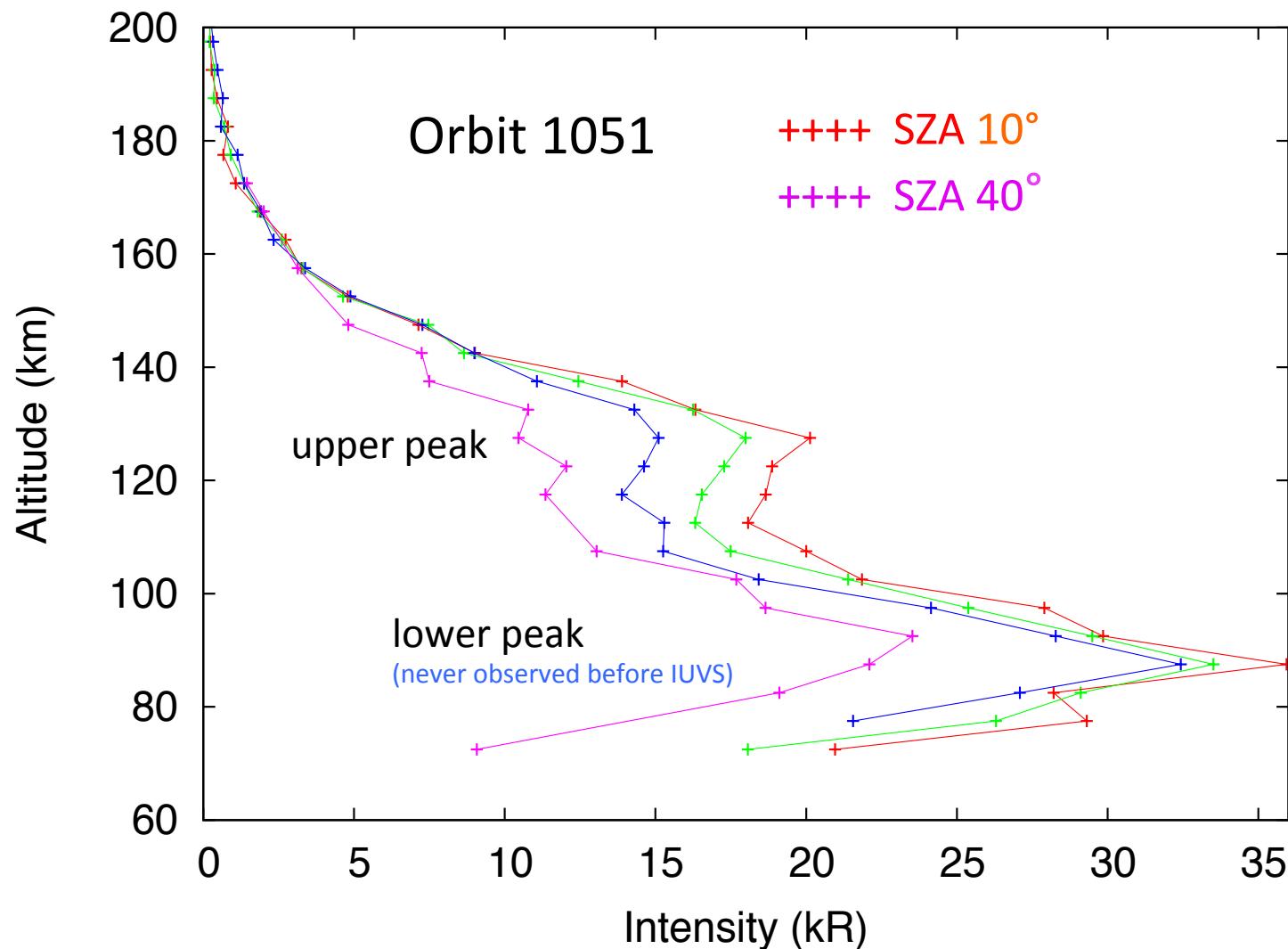
85 km

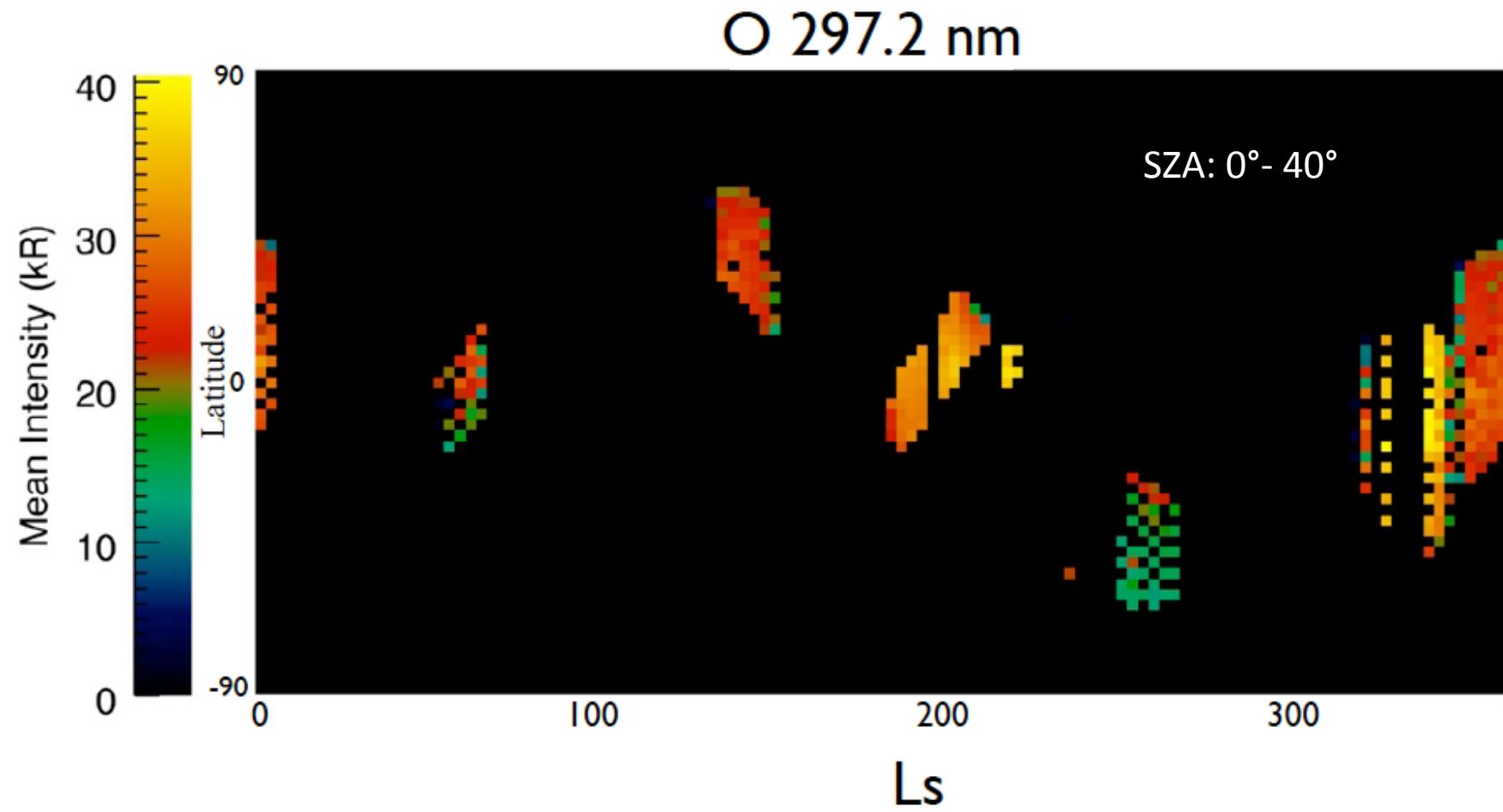
Oxygen energy levels ($2p^4$ configuration)



Both transitions to 1D and 3P levels are forbidden

Limb scans at 297.2 nm along one MAVEN orbit



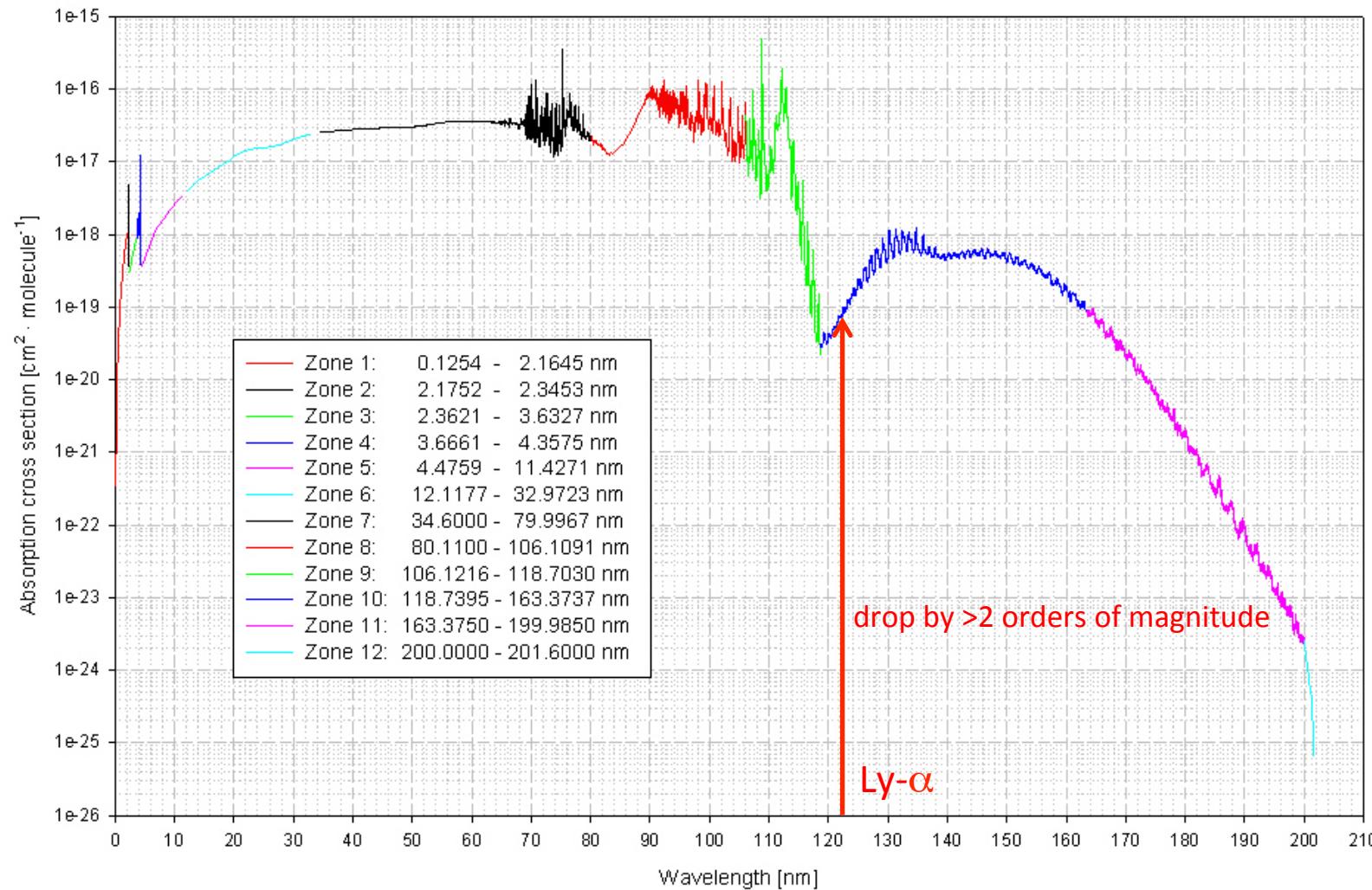


Sources of O(¹S) atoms

Process	quantity	quantum yield	number
$\text{CO}_2 + \text{photons} \rightarrow \text{CO} + \text{O}(\text{^1S})$	EUV flux	QY_1	(1)
$\text{O}_2^+ + e_{\text{th}} \rightarrow \text{O} + \text{O}(\text{^1S})$	$\text{Ly}-\alpha$ flux $\alpha_1,$	$QY_{\text{Ly}-\alpha}$ QY_2	(2)
$e_{\text{pe}} + \text{CO}_2 \rightarrow \text{CO} + \text{O}(\text{^1S})$	σ_1		(3)
$e_{\text{pe}} + \text{O}(\text{^3P}) \rightarrow \text{O}(\text{^1S}) + e_{\text{pe}}$	σ_2		(4)
$e_{\text{pe}} + \text{CO} \rightarrow \text{O}(\text{^1S}) + \text{C} + e_{\text{pe}}$	σ_3		(5)
$\text{CO}_2^+ + e_{\text{th}} \rightarrow \text{CO} + \text{O}(\text{^1S})$	α_2	QY_3	(6)

Losses: radiation + collisional quenching by CO₂, O and CO

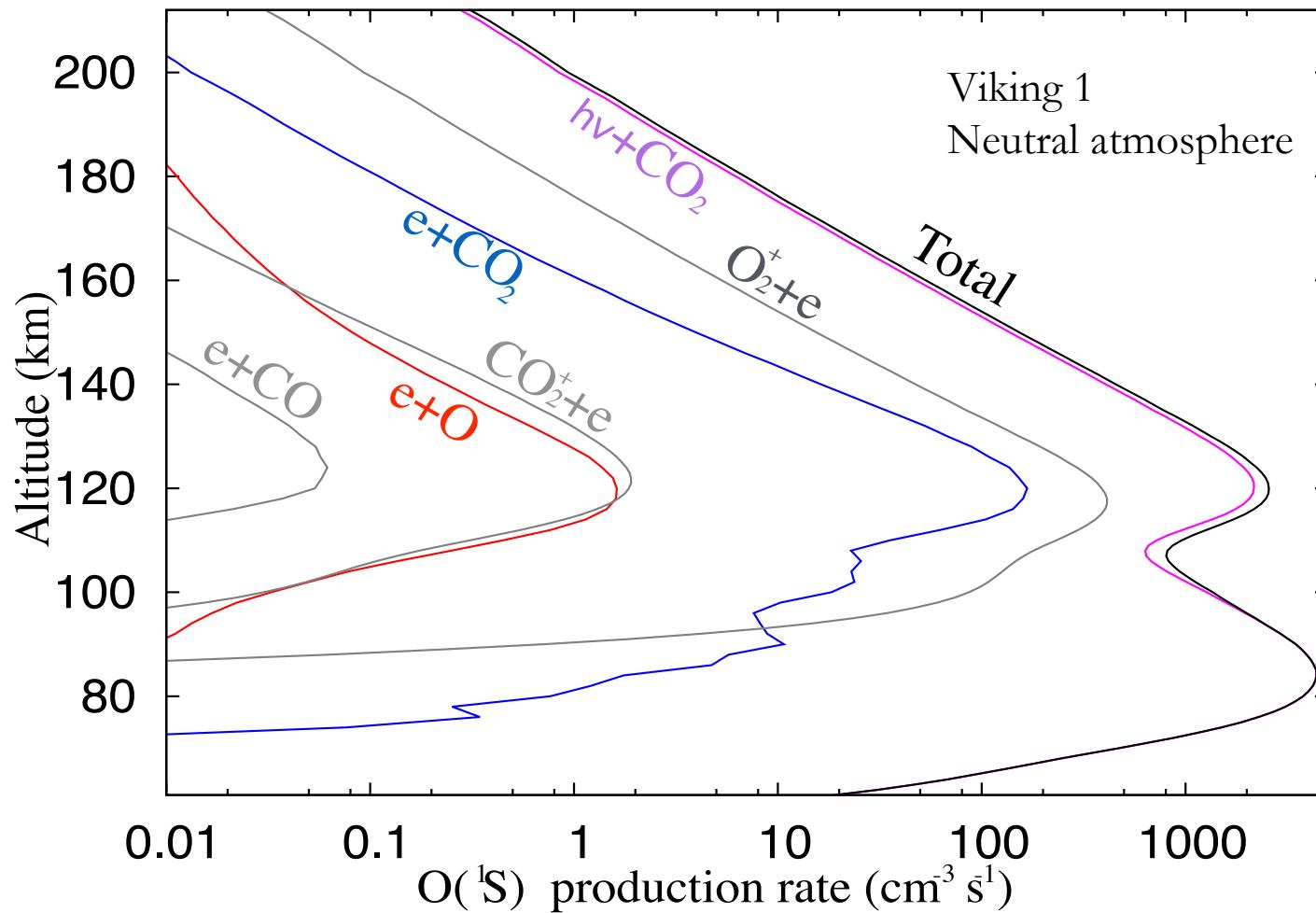
CO₂ absorption cross section (cm²)



Critically evaluated VUV absorption cross sections of carbon dioxide CO₂ at 300 K,

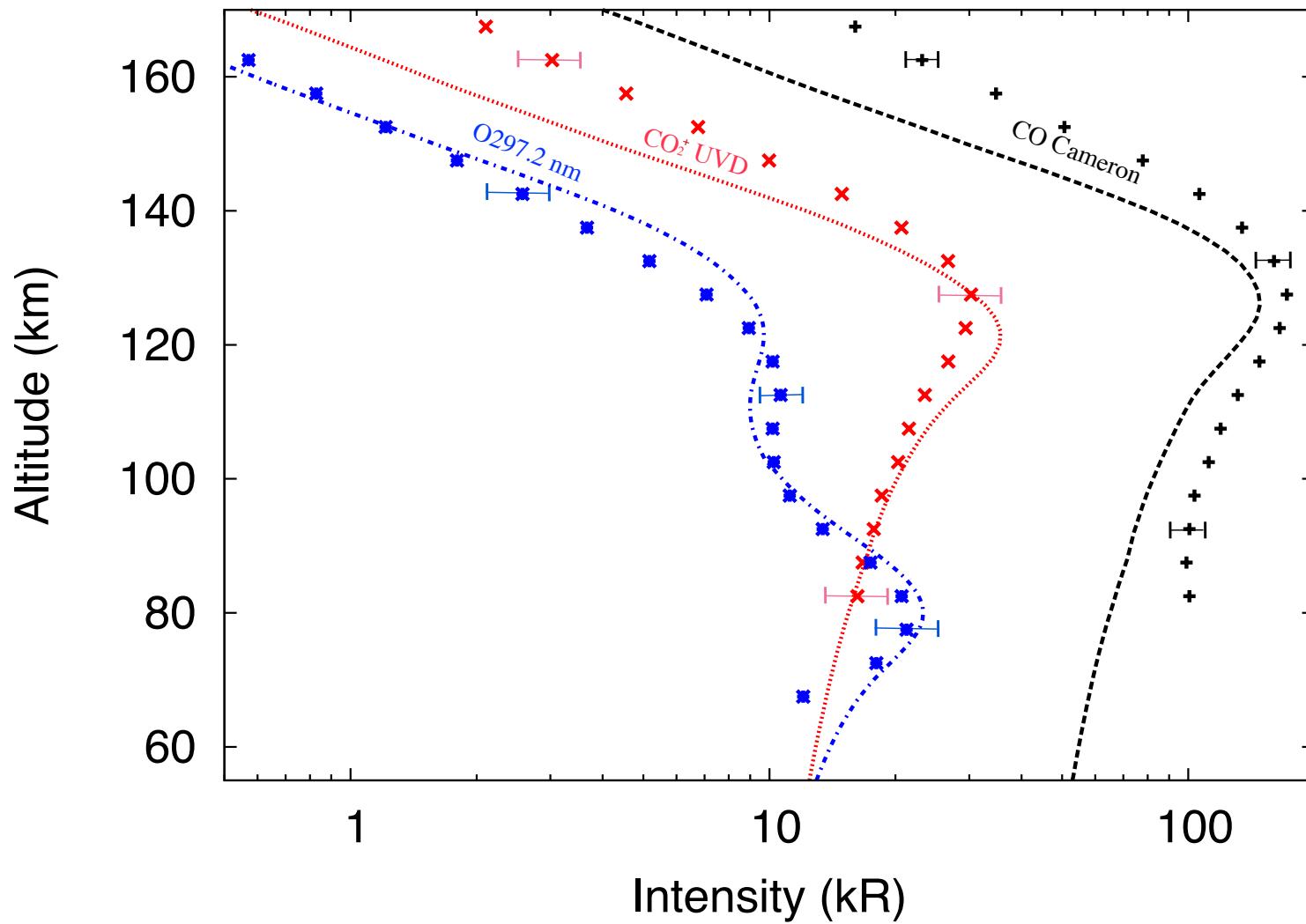
Huestis and Berkowitz, Advances in Geosciences Vol. 25:
Planetary Science (2010) 229-242 (World Scientific company).

Sources of O(¹S) atoms



Calculated from combined photochemistry and photoelectron impact model (Monte Carlo)

Other airglow emissions



CONCLUSIONS

- Limb profiles of OI 297.2 nm dayglow observed with IUVS/MAVEN show seasonal and latitudinal variations of the intensity and altitude of both peaks
- Production of O(¹S) atoms in the Martian upper atmosphere is dominated by photodissociation of CO₂ at all altitudes below 200 km
- The lower peak is produced by penetration of Ly- α solar radiation down to the 80 km region
- The quantum yield for O(¹S) production by CO₂ dissociation by Lyman- α is about 10%
- The 297.2 nm emission changing peak altitude reflects variations of the CO₂ column density and thus the pressure level